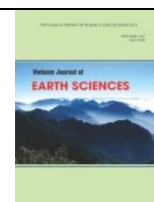




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## Paleomagnetism of Permian-Triassic volcanic sequences from Son La province, northwest Vietnam

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### ABSTRACT

Nineteen sites with 198 oriented-core samples have been collected from the Upper Permian-Lower Triassic volcanic rocks of Vien Nam Formation at Quynh Nhai locality, Son La Province, northwestern Vietnam. The characteristic remanent magnetization components carried by magnetite and hematite were successfully isolated from secondary components reveal a mean paleomagnetic direction  $D_s = 48.3^\circ$ ,  $I_s = -10.0^\circ$ ,  $\alpha_{95} = 8.0^\circ$ , corresponding to a virtual geomagnetic pole located at  $\lambda = 35.7^\circ\text{N}$ ,  $\phi = 217.4^\circ\text{E}$  and a paleo-latitude of study area situated at  $5.1^\circ\text{S}$  during the Permian time. Compared with the Late Permian-Early Triassic pole of the South China Block (SCB), the data show that crustal elements of NW Vietnam have been close to, but not unequivocally a coherent part of the SCB, since the Late Permian. Development of the parallel NW-SE striking Song Ma and Song Chay orogenic belts did not involve the closure of wide ( $> 500$  km) ocean basins.

*Key words:* Paleomagnetism; characteristic remanent magnetization; virtual geomagnetic pole; paleo-latitude; volcanic rocks.

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### 1. Introduction

The Mesozoic dispersal of the Gondwana continents and closure of Paleo-, Meso- and Neo-Tethys involved the accretion of numerous continental fragments to southern Eurasia with attending magmatism. Accurate paleo-latitude and paleo-orientation information on these fragments that now comprise Indochina are required to better understand their original configuration

relative to the South China block and timing and style of deformation associated with the Indochina block.

On the basis of biogeographic and paleomagnetic data, paleogeographic reconstruction models of Permian eastern Pangea and Tethys have been proposed by different workers (Sengor; 1979; 1984; 1987; 1989; Hutchison, 1989; Gatinsky and Hutchison, 1986; Metcalfe, 1988; 1991; 1996; 2002; 2011; Scotese et al., 1992; 1995; Dercourt et al., 1993; Ziegler et al., 1997;

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Zonenshain et al., 1985; Li et al., 1993; Li and Powell, 2001). All of the reconstruction models broadly agree on the reconstruction of the principal components of Pangea (Fig. 1).

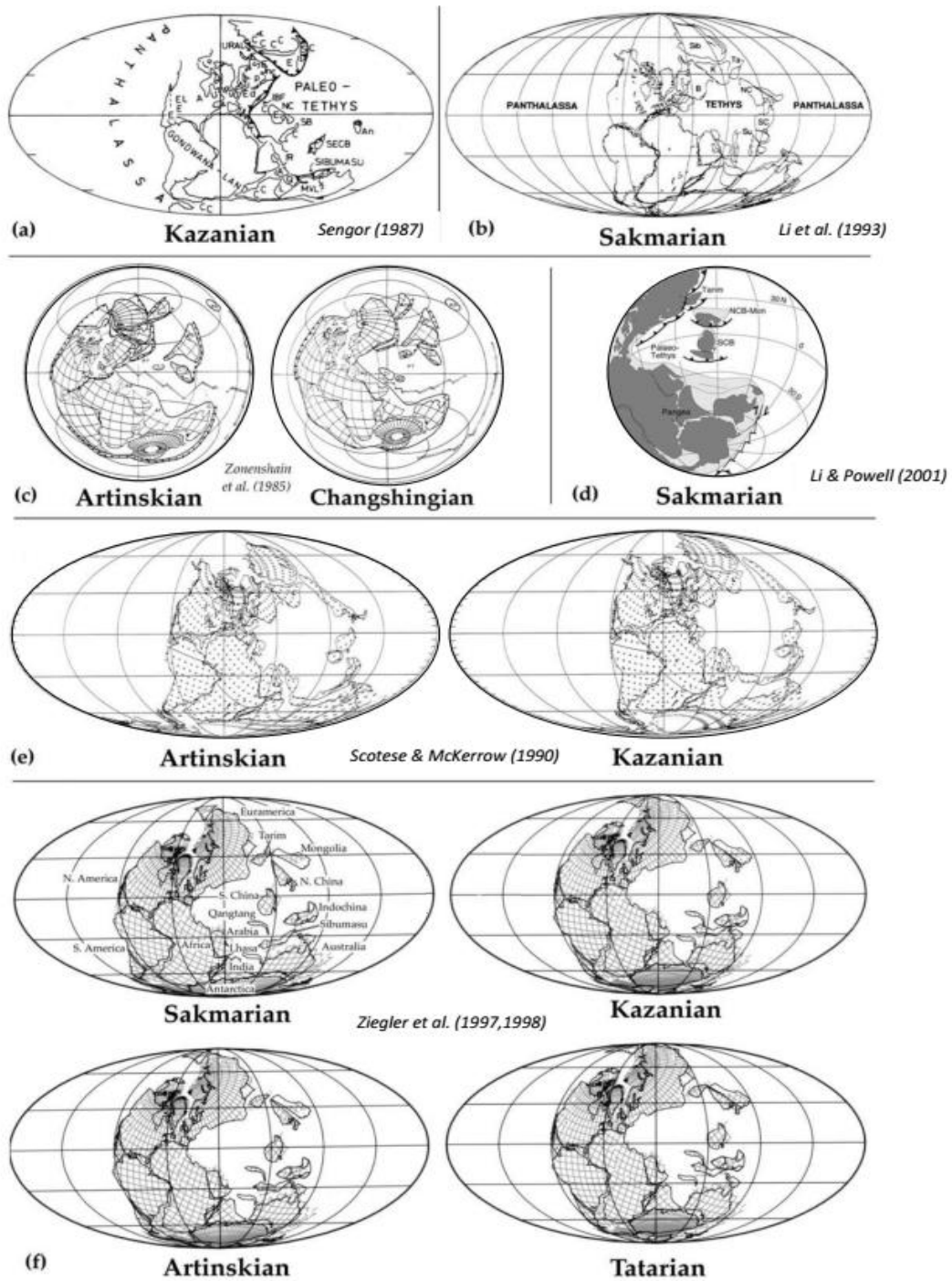


Figure 1. Paleogeographic reconstruction models of Permian eastern Pangea and Tethys (Modified from Metcalfe, 2002)

Most models for the position of South China during the Permian have the sub-continent situated on, or slightly south of the equator and slightly counter-clockwise rotated; this configuration has been supported by paleomagnetic, biogeographic and climatic data and it did not experience significant latitudinal displacement during the Permian. The position of the Indochina block, however, is somewhat in dispute. Scotese and McKerrow (1990) and Metcalfe (1998; 2002) and Li and Powell (2001) position Indochina and South China together, welded along the Song Ma Suture Zone during Late Devonian-Early Carboniferous and forming the South China- Indochina Super-terrane during the Permian, characterized by Tethyan fauna and Cathaysian flora (Hutchinson, 1989; Metcalfe, 2002). Ziegler et al. (1997; 1998), on the other hand, argue that these terranes remain separated during the Permian.

We have initiated a collection of well-exposed sequences of Upper Permian-Lower Triassic volcanic rocks exposed near the Song Da lineament in northwest Vietnam. In this paper, we report the new paleomagnetic results of Upper Permian-Lower Triassic volcanic formations from Quynh Nhai locality, Son La Province in northwestern Vietnam.

## 2. Geologic-tectonic setting and sampling area

Upper Permian-Lower Triassic volcanic rocks are widely exposed in northwest Vietnam, mainly within the Song Da rift basin located NE of the Song Ma Suture Zone that forms the boundary between the South China Block and the Indochina Block. Permian basalts within the Song Da rift are considered distal equivalents of the Emeishan rift basalts; the geochemistry of these basalts and associated Permian-Triassic komatiites suggests a plume related source (Hanski et al., 2004; Zhou et al., 2008). On the other hand, Lepvrier et al. (2004; 2008) suggest that the Permian basalts exposed north of the Song Ma Suture Zone (within Song Da rift basin) could

have been emplaced in a back-arc continental setting due to a north-dipping oceanic subduction system of Indochina beneath South China, yet an inferred calc-alkaline magmatic arc of that age is not well documented. Some authors suggest that Indochina and South China were separated by oceanic crust in the Late Permian to Early-Middle Triassic and that they collided during the Indosinian Orogeny in the Late Triassic (Zhang et al., 2006; Zhang and Cai, 2009; Cai and Zhang, 2009).

Permian flood basalts and associated mafic-ultramafic intrusions form a narrow NW-trending belt in the Song Da rift terrane. The belt is bounded by the Red River fault zone to the northeast and the Song Ma suture to the southwest (Fig. 2).

The flood basalts unconformably overlie the early Permian limestones and are concordantly overlain by the early Triassic limestone and shale (Glotov et al., 2001). Folded Triassic sedimentary rocks are unconformably overlain by Cretaceous conglomerates, sandstones and pelites (Lacassin et al., 1998).

Upper Permian - Lower Triassic volcanic rocks in the Song Da rift terrane have been divided into the Cam Thuy and Vien Nam formations. The Upper Permian Cam Thuy Formation consists of mafic extrusive rocks of tholeiitic affinity that are widely exposed in the southeast part of the Song Ma anticline within Thanh Hoa, Ninh Binh, and Hoa Binh, Son La to the north of Song Ma suture. The volcanic rocks of the Cam Thuy Formation are, in general, basalt, basaltic andesite, or andesite and are up to 1000 m thick (Hoa Binh area). They are dominated by high magnesium, low titanium and alkaline basalt, equivalent to tholeiite series rocks (Polyakov et al., 1999; Tran Trong Hoa, 2001).

The Vien Nam Formation (Phan Cu Tien, 1977) is widely distributed within the Song Da rift terrane, from the Vien Nam area - Ba Vi extending through Kim Boi (Hoa Binh) to

Van Yen, Ta Khoa and Nam So (Son La) areas in northwest Vietnam. Tran Trong Hoa (2001) describes the extrusive rocks of Vien Nam Formation as belonging to a combination of basalt (trachybasalt) - trachyandesite - trachydacite, which are sub-alkaline basalts with high Ti and high alkaline compositions. Based on isotopic age determinations

(Polyakov et al., 1999; Tran Trong Hoa, 1996; Hanski et al., 2004), the interpreted age of the basaltic rocks is between about  $257 \pm 24$  Ma and  $270 \pm 15$  Ma, corresponding to the Late Permian-Early Triassic, and is inferred to be coeval with the Emeishan basalts exposed in South China.

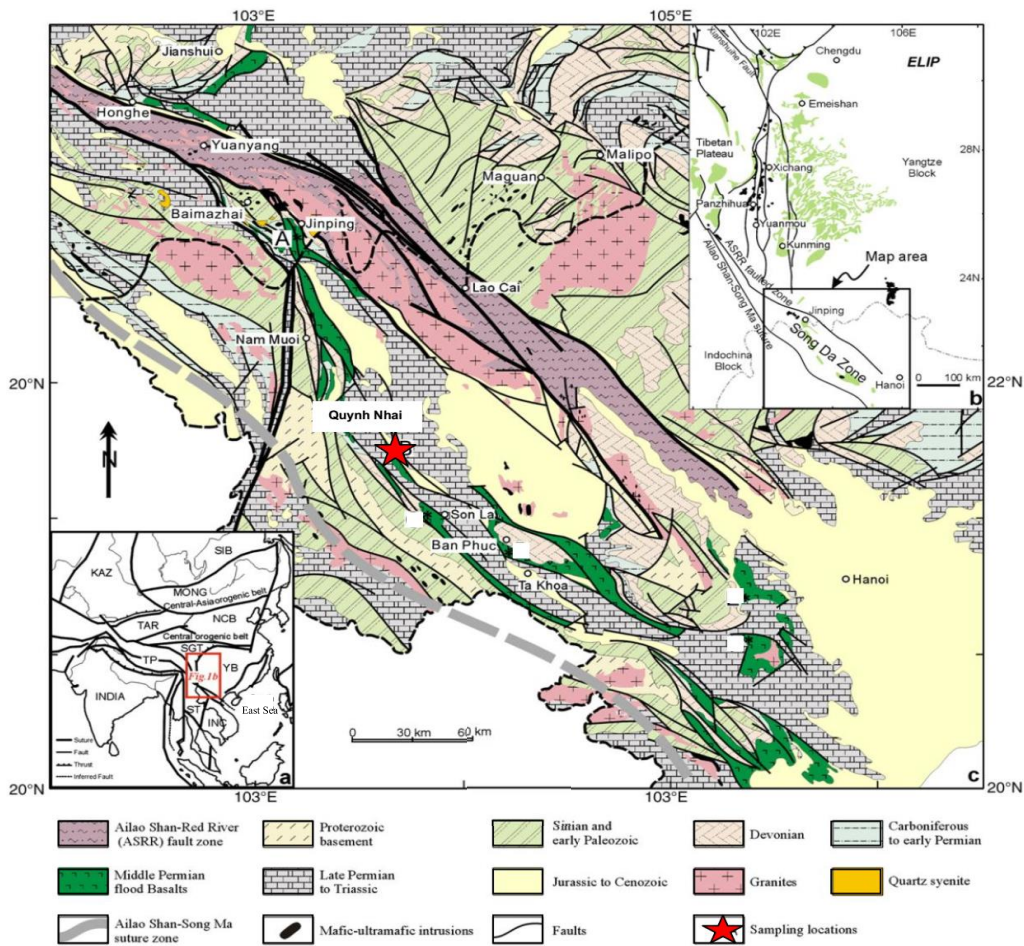


Figure 2. Geological sketch map of the NW Vietnam region showing the study area, modified from Wang et al. (2007). The major blocks shown in inset (a) are: SIB, Siberia; MONG, Mongolia; NCB, North China Block; YB, Yangtze Block; INC, Indochina Block; TAR, Tarim; KAZ, Kazakhstan; TP, Tibetan Plateau; SGT, Songpan-Ganze terrane. The Emeishan Large Igneous Province in SW China and northern Vietnam is shown in inset (b)

At Quynh Nhai locality, there are several outcrops of vesicular basalt and basaltic andesite of Vien Nam Formation on the road from Quynh Nhai town to the Pa Uon Bridge

crossing the Da River. These basalts had been periodically emplaced at the surface revealed by at least 8-10 flows of ~0.5 m in thickness with flow orientations varying from 192-

218°/44-79° SW that enables to have a reliable tectonic control for these basalts (Fig. 3). Nineteen sites with 198 oriented-core samples have been collected from these outcrops.



Figure 3. Photo showing the outcrop of the lower Triassic basalt flows of the Vien Nam formation at Quynh Nhai locality, Son La, NW Vietnam

### 3. Rock magnetism

Individual specimens of 25 mm in diameter and 22 mm in length have been prepared from each sample in the laboratory. Natural remanent magnetizations (NRMs) were measured by 2-G Enterprise's cryogenic magnetometer. One specimen from each sample was subjected to stepwise thermal demagnetization up to 690°C using a TDS-1 thermal demagnetizer with a residual field of less than 5 nT. Demagnetization results for each specimen were plotted on orthogonal vector diagrams (Zijderveld, 1967) to assess the component structure as well as on equal-area projections to evaluate directional stability. Principal component analysis (Kirschvink, 1980) was used to estimate component's directions. Site-mean directions were calculated using Fisherian statistics (Fisher, 1953). The natural remanent magnetizations of basaltic rocks are rather

strong, ranging from 15.8 mA/m to 3.78 A/m, and consist of two or more components (Figs. 4, 5). These rocks show excellent response in the progressive alternating field and/or thermal demagnetization. For any single section sampled, there is a mix of carriers of the characteristic remanent magnetization (ChRM), with some flows having low Ti magnetite as the principal carrier of the ChRM and other flows having hematite as the ChRM. The experiments of rock's magnetic susceptibility during the heating and cooling processes shown in Fig. 6 also revealing the behavior of ChRM multi-components. The magnetic minerals in the rock structure are shown in Fig. 7. Rock magnetic data and petrography support the retention of an early-acquired thermo-remanent magnetization; anisotropy of magnetic susceptibility (AMS) data is not consistent with a pervasive post-emplacement deformation fabric (Fig. 8).

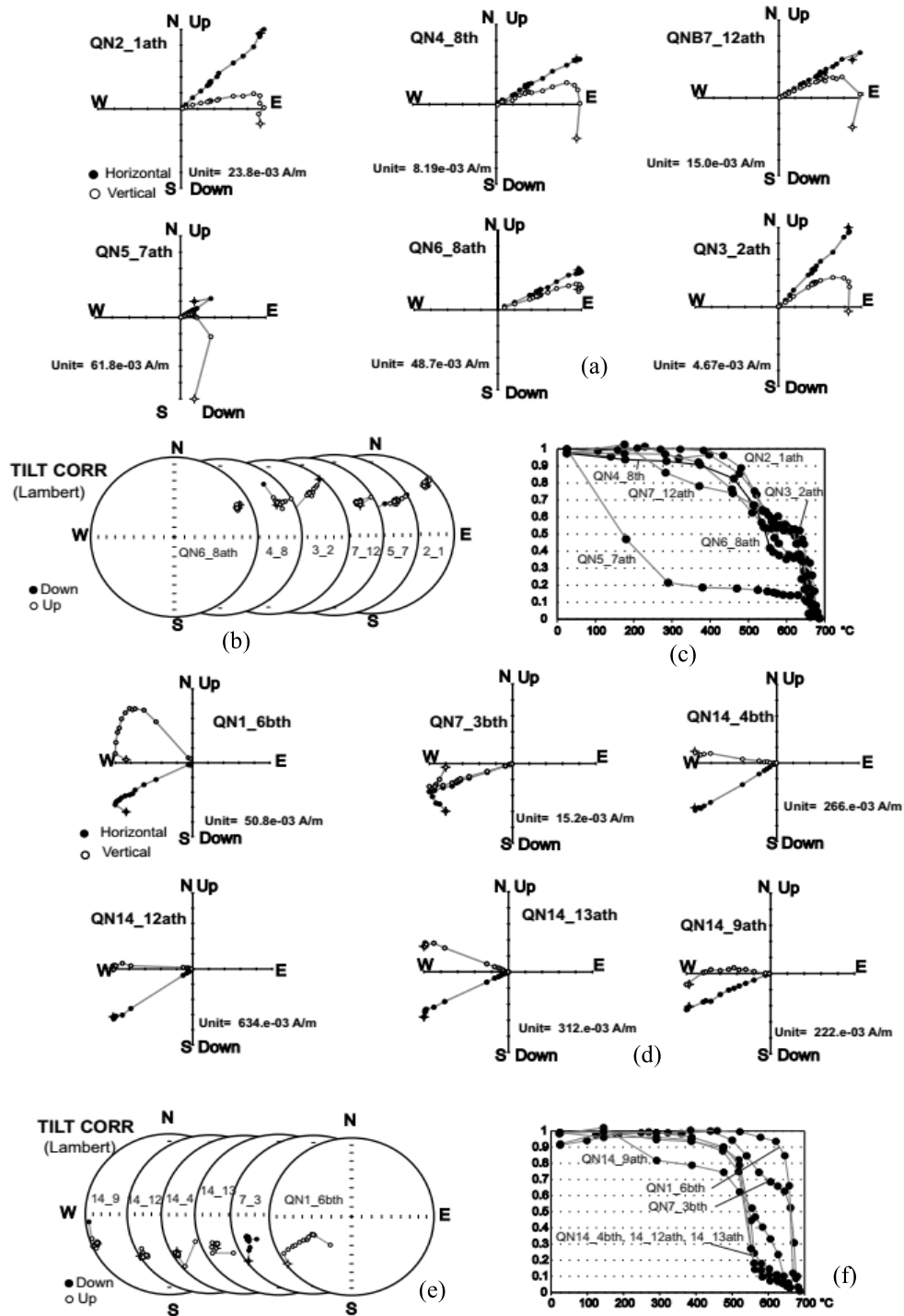


Figure 4. Representative thermal demagnetization plots of rock specimens; (a) & (d): Zijderveld plots, (b) & (e): Lambert equal-area plots, (c) & (f): normalized NRM plots

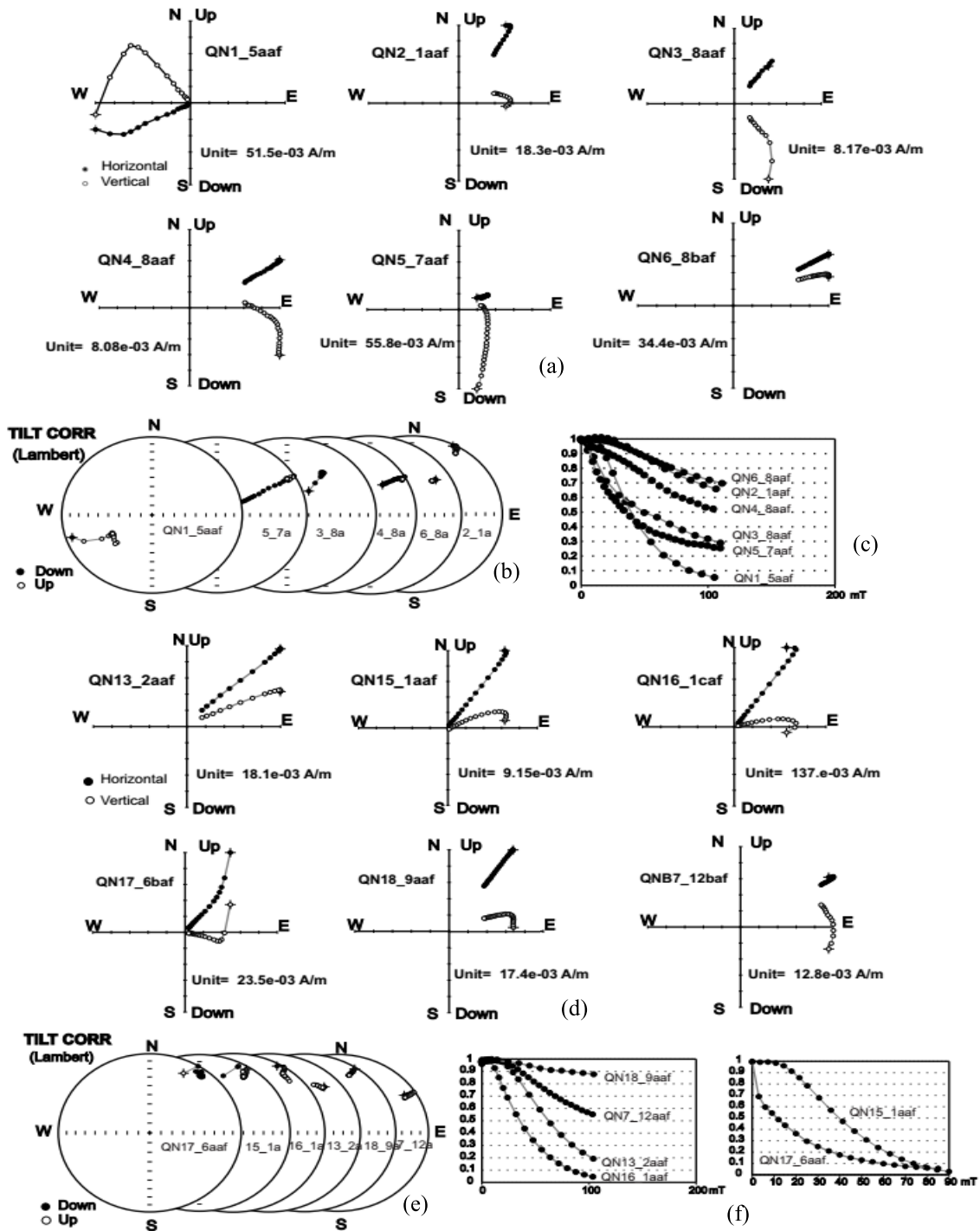


Figure 5. Representative AF demagnetization plots of rock specimens; (a) & (d): Zijderveld plots, (b) & (e): Lambert equal-area plots, (c) & (f): normalized NRM plots

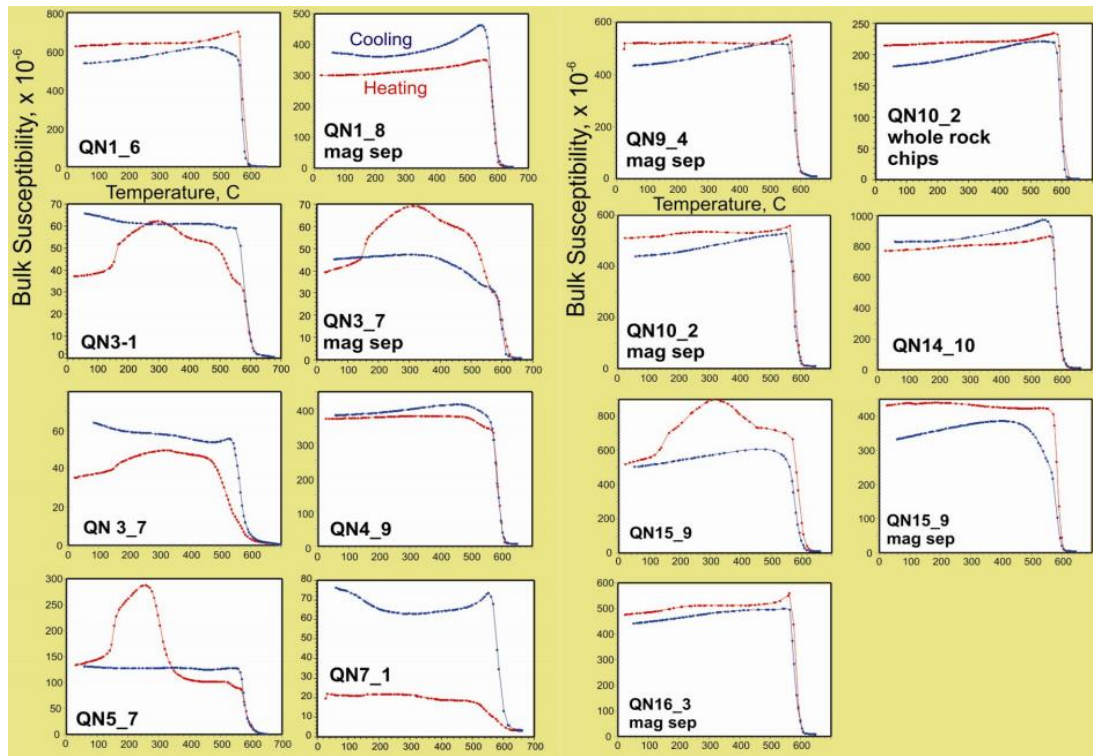


Figure 6. Bulk magnetic susceptibility of representative rock specimens during heating and cooling process showing the behavior of ChRM multi-components of low Ti magnetite and hematite

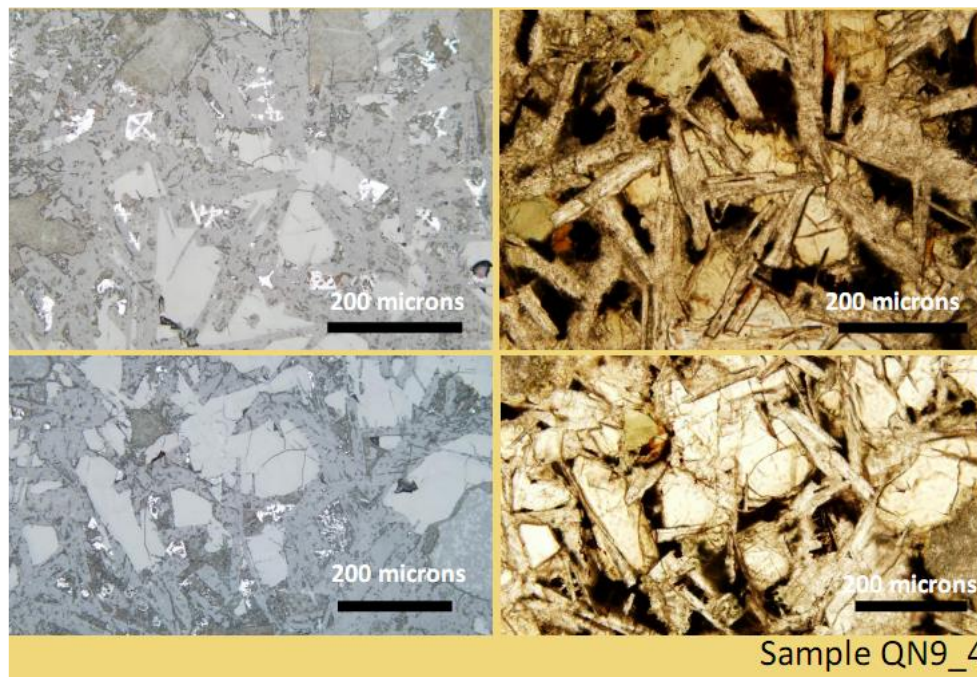


Figure 7. Microscopic photos show the structure of magnetic-bearing minerals



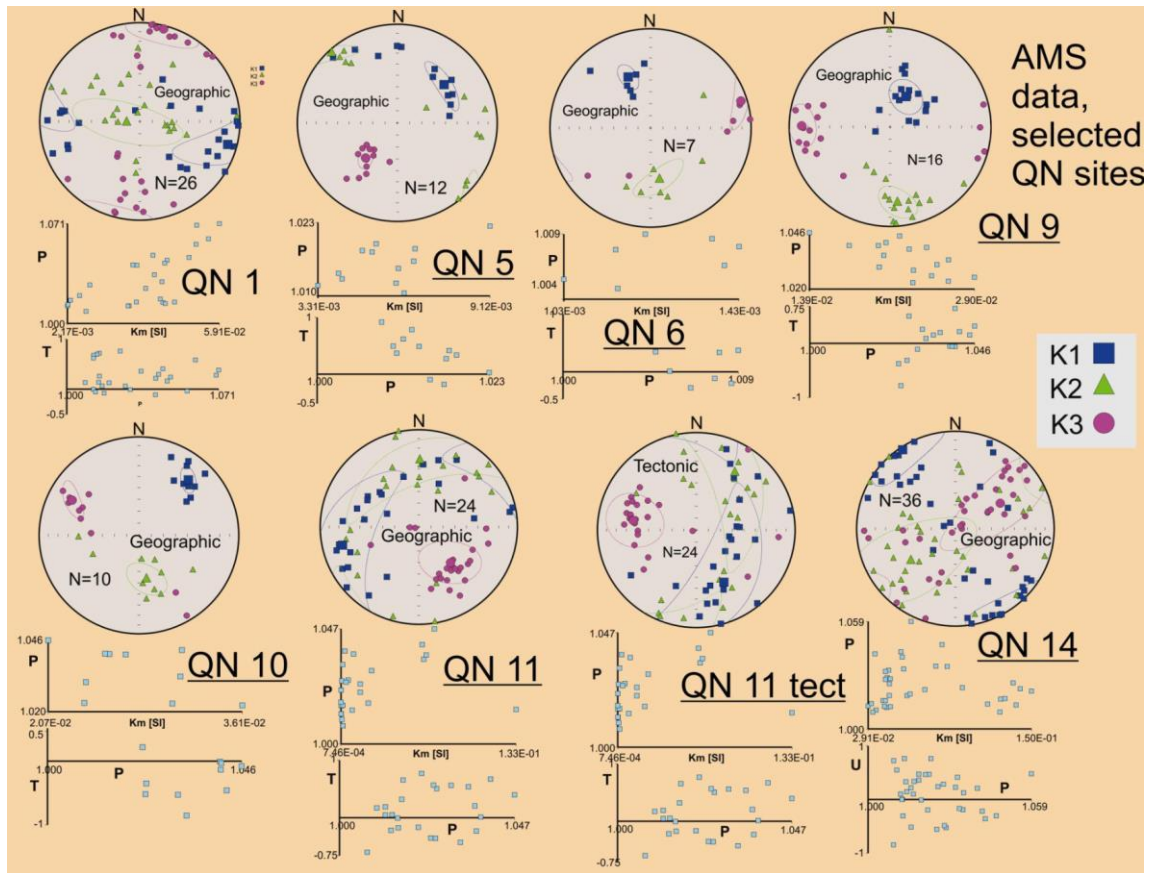


Figure 8. Anisotropy of magnetic susceptibility (AMS) of selected samples

#### 4. Paleomagnetic results

The site-mean directions of ChRM components are plotted on the Lambert equal-area projection (Fig. 9) and their values are listed in Table 1. At the site level, magnetization directions are consistent and estimates of mean directions have 95% confidence estimates of less than 7°. Preliminary between-site data suggest that a reasonable magnitude of paleosecular variation has been recorded. In situ estimated site mean directions do not resemble predicted time-averaged geomagnetic field directions for the study area for Triassic and younger time, using paleomagnetic poles from the South China Block and therefore we argue that these volcanic sections have not been

pervasively remagnetized. Magnetizations are of dual polarity with inclinations generally less than 20°.

The formation mean direction of 13 sites yields the values  $D_s = 48.3^\circ$ ,  $I_s = -10.0^\circ$ ,  $\alpha_{95} = 8.0^\circ$ ,  $k = 27.7$  ( $k_s/k_g = 1.37$ ); corresponding to a virtual geomagnetic pole (VGP) located at  $\lambda = 35.7^\circ\text{N}$ ,  $\phi = 217.4^\circ\text{E}$ . The paleopole position of Quynh Nhai basaltic rocks is basically indistinguishable from the paleopole of Late Permian-Early Triassic paleopole for the South China Block (Fig. 10). The similar paleopole positions of two distal basaltic rock formations from NW Vietnam and South China further corroborate the hypothesis of coeval time emplacement and originated from the same source for these formations. The

paleolatitude of the study area calculated from the formation mean direction is  $\lambda_{obs} = -5.1^\circ N$  (or  $5.3^\circ S$ ) indicating that the Quynh Nhai basaltic rock formation was emplaced on the Earth during the Late Permian-Early Triassic time near the equatorial region.

Using the mean paleopole of Upper Permian - Early Triassic volcanic rock formations from South China Block (Table 2), the expected paleomagnetic direction of Quynh Nhai locality is computed and has the values:  $D_{ex} = 33.6^\circ$ ,  $I_{ex} = -10.3^\circ$ ,  $\lambda_{ex} = -5.2^\circ N$ .

Table 1. Paleomagnetic results of Late Permian-Early Triassic volcanic rocks from Quynh Nhai locality

Site	Location		Az/Dp (o)	n/N	ChRM						VGP		dp (o)	dm (o)	$\lambda_{paleo}$ ( $^\circ N$ )
	$\lambda(^{\circ}N)$	$\phi(^{\circ}E)$			Dg	Ig	Ds	Is	$\alpha_{95}$	k	$\lambda(^{\circ}N)$	$\phi(^{\circ}E)$			
QN01*	21.524	103.733	198/60	12/12	229.6	4.6	244.6	-44.1	6.6	44.0	-31.3	355.8	5.2	8.3	-25.8
QN02	21.680	103.616	198/60	5/5	65.3	-58.5	40.8	-6.9	12.1	41.0	42.9	220.9	6.1	12.2	-3.5
QN03	21.680	103.616	198/60	8/8	66.1	-61.5	39.1	-9.4	14.7	15.1	43.5	223.5	7.5	14.9	-4.7
QN04	21.680	103.616	198/60	11/11	83.2	-40.6	61.7	-2.8	6.9	44.4	25.6	206.3	3.5	6.9	-1.4
QN05	21.680	103.616	198/60	10/10	79.4	-44.9	56.5	-3.4	4.2	135.8	30.1	209.1	2.1	4.2	-1.7
QN06	21.680	103.616	198/60	5/5	100.7	-48.9	61.2	-17.8	3.3	533.7	22.6	214.1	1.8	3.4	-9.1
QN07	21.680	103.616	198/60	6/6	101.7	-47.0	63.2	-17.1	5.2	164.2	21.0	218.8	2.8	5.4	-8.8
QN7B	21.680	103.616	198/60	8/8	102.1	-48.7	61.8	-18.4	5.0	122.6	21.9	214.1	2.7	5.2	-9.5
QN08*	21.665	103.628	198/60	12/12	339.1	54.6	235.2	52.9	10.2	19.2	13.8	238.8	9.7	14.0	-33.5
QN09*	21.669	103.633	212/44	12/12	171.0	-62.1	72.5	-61.7	2.0	454.8	-2.7	239.3	2.4	3.2	-42.9
QN10*	21.669	103.633	212/44	9/9	182.3	-48.8	103.2	-69.9	3.7	193.4	-25.0	244.2	5.5	6.4	-53.7
QN13	21.670	103.629	218/79	6/6	241.7	-78.9	33.2	-21.1	8.3	66.2	43.9	235.4	4.6	8.7	-10.9
QN14	21.671	103.626	218/79	14/14	271.3	59.8	242.0	-7.5	3.8	112.6	27.3	200.8	1.9	3.8	3.8
QN15	21.679	103.617	192/56	9/9	64.5	-59.0	36.8	-12.7	4.9	112.3	44.3	227.4	2.5	5.0	-6.4
QN16	21.679	103.617	192/56	10/10	65.1	-58.8	37.1	-12.7	3.8	159.5	44.0	227.1	2.0	3.9	-6.5
QN17	21.679	103.617	192/56	13/13	49.6	-42.8	38.7	5.8	3.2	164.6	48.0	214.7	1.6	3.3	2.9
QN18	21.679	103.617	192/56	7/7	70.6	-64.2	35.0	-18.3	2.8	464.4	43.7	232.1	1.5	2.9	-9.4
Mean G	21.678	103.618		13/17	79.6	-57.5			9.5	20.2	5.5	52.6	10.1	13.8	38.1
Mean S	21.678	103.618		13/17			48.3	-10.0	8.0	27.7	35.7	217.4	4.1	8.1	-5.1

Note: (\*)- sites are excluded from the calculation of site mean; Az/Dp: Azimuth of dip/Dip; ChRM: characteristic remanent magnetization; VGP: Virtual geomagnetic pole;  $\lambda_{paleo}$ : Paleolatitude of the site

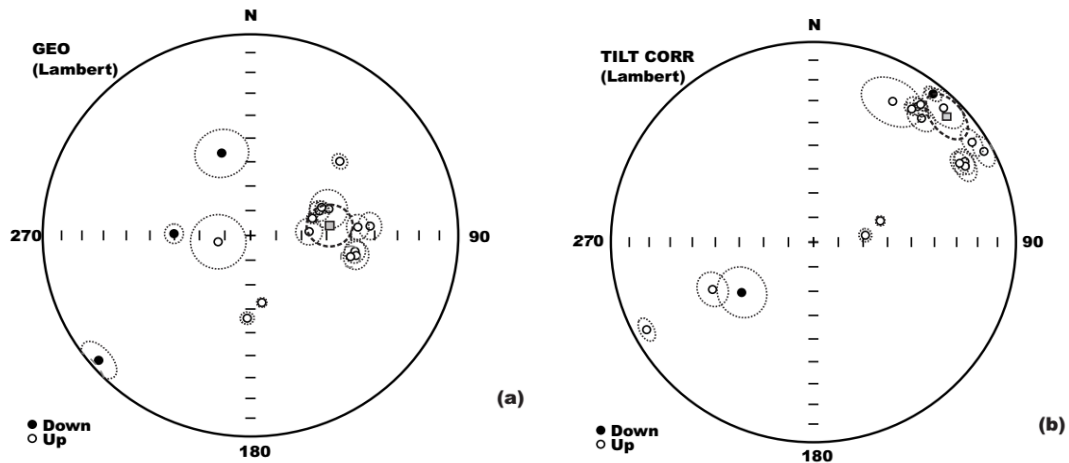


Figure 9. Lambert equal-area diagram showing the site mean directions with the ellipse of 95% confidence. (a) in geographic coordinates; (b) in stratigraphic coordinates

Table 2. Late Permian - Early Triassic virtual geomagnetic poles from South China

Age	Locality		VGP coord.		A <sub>95</sub>	Rock type	Reference
	λ	φ	λ	φ			
South China Block:							
P2	29.6	103.4	52.7	252.1	6.0	Basalt and red sandstone (Emeishan, Sichuan)	McEnhinny et al., 1981
P2	29.6	103.4	49.7	252.0	2.8	Emeishan basalt	Zhao and Coe, 1989
P2	29.6	103.4	54.1	241.8	19.0	Basalt and limestone (Emeishan, Sichuan)	Chan et al., 1984
P2	29.6	103.5	38.5	231.6	4.9	Emeishan basalt, Xihekou	Liu et al., 1985
P2	28.1	102.9	58.9	246.1	6.6	Basalt, Zhaojue, Sichuan	Liu et al., 1985
P2	27.6	101.9	63.5	264.3	9.8	Basalt, Yanguan, Sichuan	Liu et al., 1985
P2	28.3	103.0	54.3	251.2	20.0	Basalt, Meigou, Sichuan	Zhou et al., 1986
P2	26.8	101.8	25.6	216.4	12.6	Basalt, Miyi, Sichuan	Zhou et al., 1986
P2	26.7	102.9	53.5	241.8	10.1	Basalt, Huidong, Sichuan	Huang et al., 1986
P2	26.1	103.1	52.5	226.0	25.0	Basalt, Dongchuan, Yunnan	Huang et al., 1986
P2	25.9	100.6	24.7	204.3	24.8	Basalt, Binhchuan, Yunnan	Huang et al., 1986
P2	26.4	105.7	29.3	235.3	13.4	Basalt, Xiongjiachang, Guizhou	Lin, 1984
P2	25.6	103.0	50.0	241.0	6.1	Basalt, Kunming, Yunnan	Fang, Vander Voo, 1990
P2	26.5	101.8	23.6	210.3	9.1		Zhou et al., 1988
P2	26.5	101.8	28.9	218.2	6.4		Zhou et al., 1988
P2	28.3	103.0	57.0	277.1	12.7		Zhou et al., 1988
P2	28.3	103.0	47.1	231.0	19.0		Zhou et al., 1988
P2	29.6	103.4	54.1	241.8	19.5		Lung 1984
P2	29.6	103.4	51.2	232.5	2.9		Zhang 1984
P2	29.6	103.4	48.9	251.1	4.0		Zhao & Coe 1987
P2	26.4	104.7	41.5	222.7	4.5		Huang, Opdyke, 1998
P2	30.1	102.7	51.7	254.1	5.1		Zhuang et al., 1989
P2-T1	32.4	105.5	47.2	226.3	4.2	Red limestone, Guanyuan, Sichuan	Steiner et al., 1989
P2-T1	32.4	105.5	47.9	225.1	5.4	Red limestone, Guanyuan, Sichuan	Heller et al., 1988
P2-T1	30.0	106.3	38.5	209.8	3.3	Red limestone, Hechuan, Sichuan	Steiner et al., 1989
P2-T1	29.6	105.6	45.8	225.2	7.0	Red limestone, Chongqing, Sichuan	Steiner et al., 1989
P2-T1	30.0	107.0	39.8	216.2	5.3	Red limestone, Changshou, Sichuan	Enkin, 1990
P3-T1	29.5	106.0	40.7	215.0	6.8	Sichuan	Enkin et al., 1992b
P3-T1	32.4	105.5	50.3	222.9	3.3		Li et al., 1988
P3-T1	31.1	119.7	51.1	230.2	-		Li & Wang, 1989
T1	31.3	110.6	-40.3	37.3	26.3		Steiner et al., 1989
T1*	29.2	117.3	55.4	354.1	2.1		Steiner et al., 1989
T1	28.6	106.9	46.3	219.3	11.0		Steiner et al., 1989
T1*	30.1	102.7	71.0	276.7	4.6		
T1	30.3	106.8	42.6	217.3	8.9		
T1	28.0	105.5	-44.0	37.4	6.8		
T1	27.4	100.0	58.6	205.9	7.9		
T1	31.2	110.5	-40.4	37.0	26.3		
T1	32.1	106.2	42.5	214.7	5.9		
Mean of 37/39 poles:			47.5	228.9	4.6		

Note: (\*) Poles are excluded from the calculation of mean pole

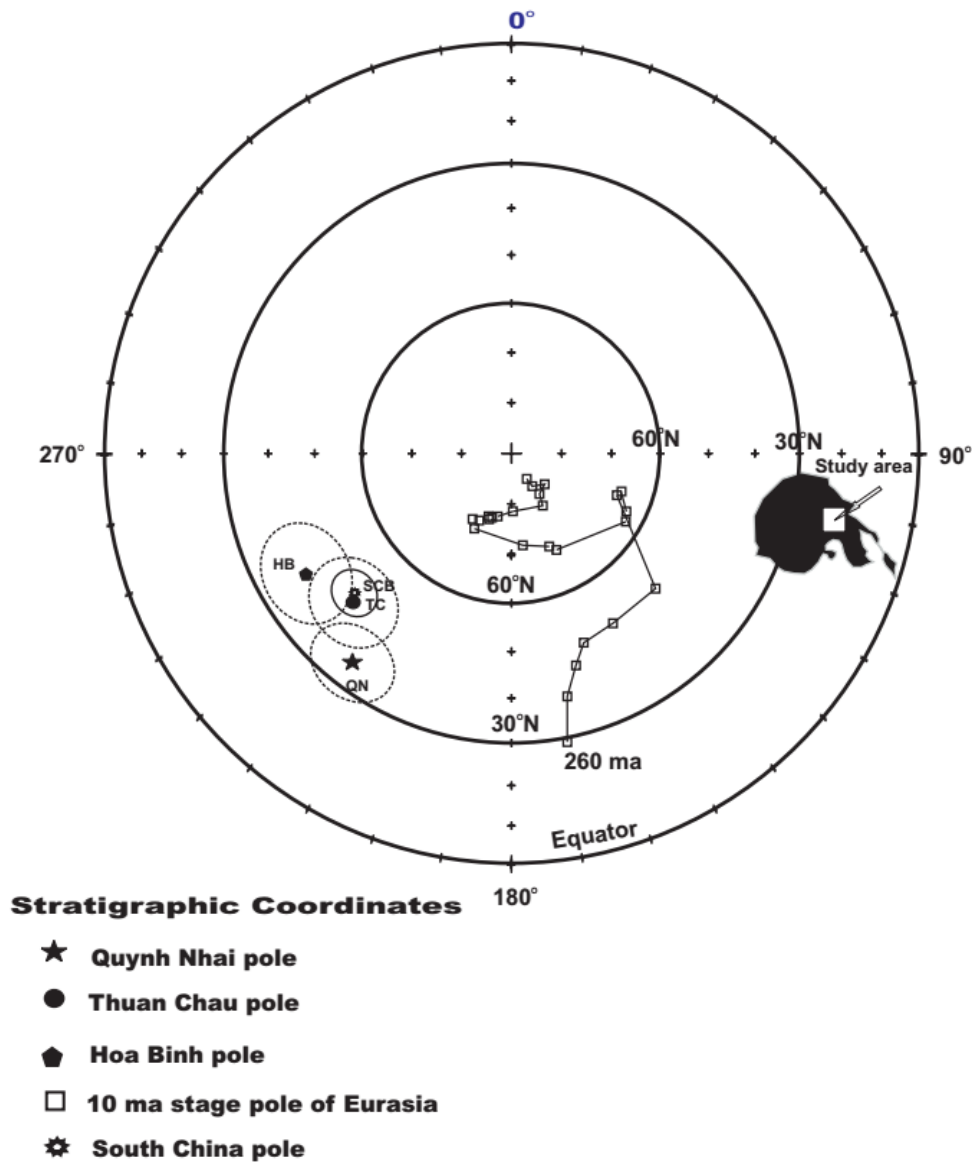


Figure 10. Equal-area diagram showing the positions of virtual geomagnetic poles (VGP) of Quynh Nhai locality relative to coeval paleopoles of Hoa Binh, Thuan Chau areas and the South China Block, and the apparent polar wander path (APWP) of Eurasia after (Yang & Besse, 2001)

## 5. Discussions

According to the paleogeographic reconstruction models, the northwestern Vietnam, in general, and the Quynh Nhai locality, in particular, belongs to the South China Block but not to the ancient Indochina Block as described by Metcalfe (2011, 2013)

because the Song Ma Suture, the boundary between two blocks, situates to the southwest of this region. The rotation and latitudinal translation of Quynh Nhai locality relative to the South China Block are calculated and have the values as follow:  $R = 14.7^{\circ} \pm 7.4^{\circ}$ ,  $d = -0.1^{\circ} \pm 7.4^{\circ}$  indicating that the Quynh Nhai

locality has been relatively stable with respect to the South China Block, in terms of latitudinal position, since the Late Permian - Early Triassic time, while has been undergoing a clockwise (CW) rotation of  $14.7^{\circ} \pm 7.4^{\circ}$  relative to the South China Block. The paleolatitude of Quynh Nhai basalts is consistent with that of Late Permian Thuan Chau basalts and Late Permian-Early Triassic Hoa Binh basalts reported by the authors in the other papers, while the observed declinations of Thuan Chau and Hoa Binh basalts showing no rotation of these areas relative to the South China Block, as well as the paleomagnetic results of Cretaceous redbeds from NW Vietnam (Cung et al., 2000; Takemoto et al., 2005) indicating that the observed CW rotation of Quynh Nhai locality might reflect a local rotation related to the folding process of these rocks.

The consistency of paleomagnetic information on relative tectonic displacement of the northwestern Vietnam region obtained from different rock formations with different geological ages shows that crustal elements of NW Vietnam have been close to, but not unequivocally a coherent part of the South China Block, since the Late Permian-Early Triassic. Development of the parallel NW-SE striking Song Ma and Song Chay orogenic belts did not involve the closure of wide (> 500 km) ocean basins in this region.

## 6. Conclusions

The consistency of paleomagnetic results of basaltic rocks from northwestern Vietnam with the paleomagnetic data of Late Permian Emeishan basalt from South China indicates that these basaltic formations might have a similar origin and coeval age.

The Quynh Nhai locality, in particular, and the northwestern Vietnam region, in general, have been close to, but not unequivocally a coherent part of the South China Block and situated near the Equator since the Late Permian-Early Triassic.

Development of the parallel NW-SE striking Song Ma and Song Chay orogenic belts did not involve the closure of wide (> 500 km) ocean basins. The Red River fault has played a minor role during the Cenozoic India-Eurasia collision and there is insignificant displacement along this fault in terms of paleomagnetism.

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