ASSESSMENT OF FLOODING IN THE DOWNSTREAM AREA OF KIEN GIANG RIVER BASIN, QUANG BINH PROVINCE BASED ON USING THE MATHEMATICAL MODEL AND REMOTE SENSING TECHNIQUE

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ABSTRACT: Floods frequently occur in the central region in general and in Quang Binh province in particular, including the Kien Giang river basin. In recent years, due to the effects of global climate change and socio-economic development, natural disasters and floods are increasing and threatening the lives of people. This affects the socio-economic development of Quang Binh province. This paper presents the results of a study on the flooding and flooding levels of some typical floods in the downstream area of Kien Giang river, Quang Binh province based on the combination of using MIKEFLOOD model with remote sensing. This is one of the modern research methods that are quite useful in simulating, calculating and forecasting natural disasters occurring in coastal estuaries including floods. Combined use of these two methods yielded, quite accurate and similar results in calculating the flood level and inundation in the Kien Giang river basin. The authors have applied the MIKEFLOOD model to simulate the continuous evolution of inundation processes combined with remote sensing technology to compare and verify the calculation results of the model at the occurrence time of typical floods from 1999 up to now in the study area. The results of this article are the basis for warning and suggesting some measures to prevent flooding before the occurrence of heavy rains in Kien Giang river basin.

Keywords: MIKE, remote sensing, shoreline separation, landsat, inundation, Kien Giang river.

INTRODUCTION

Flooding is one of the major problems in the world and in Vietnam. Flood disaster is one of the natural disasters that encompass a large area.

Floods occurring in central river basins in general and Kien Giang river basin in particular are one of the most frequent natural disasters that threaten people's lives as well as socioeconomic development in the region and Quang Binh province.

Kien Giang river is the confluence of many rivers and streams, with a basin area of $2,650 \text{ km}^2$, located in the depression of the Central Coast (fig. 1). The basin is severely

divided, so the river network develops with a river network density of 0.84 km/km². The downstream area of the river is prone to flooding during rainy season. Every year, the area is always affected by the floods, resulting in economic, political, cultural and social damages in the area and threatening Dong Hoi city [1].

In the late twentieth and early XXI centuries, natural disasters and floods have occurred continuously in Kien Giang river basin. In particular, the historic floods in November 1999 and December 1999, October 2011, October 2013 and May, June 2016 have left hundreds of dead and missing, smothering thousands of houses, crops, with estimated losses of trillions VND. Floods have caused serious consequences, thousands of households have to move from landslide or flood areas, environmental ecosystems in coastal areas of estuaries are severely degraded [2].



Fig. 1. Study area

This article introduces a method of assessing and simulating some typical flood events in Kien Giang river basin by combining mathematical model and remote sensing technology and geographic information system. The simultaneous use of these methods produces the best results in evaluating flood events compared to previous studies. The results of this study will be the basis for warning and flood prevention planning. In addition, they are a good reference for local decision-makers.

METHODOLOGY

Statistical methods

The collected data focus on meteorological, hydrologic, topographic, geological, socioeconomic data, current status of works in the basin, upstream reservoir operation process. Documentation is carried out according to process, norms (if any) and scientifically stored. The methods of synthesis, statistics, analysis, processing of existing basic data, data for research are used to calculate the study area.

The mathematical modeling method

This is a very important and effective method to implement the research topic. Based on the results obtained in the research on the basis of methodology, PhD students will select appropriate establish or modern mathematical models to calculate and simulate hydrodynamics, flood events. The object, the research problem are very complex; therefore, the tools, calculation software must be strong enough. At present, with the development of science that enables research to be very effective, is the mathematical modeling toolkit. Depending on the purpose of the study, one, two and three-dimensional numerical models will be suitable for application. Study on synthesis of flow in the catchment from rain, regulation of the lake often uses the models HEC-RESSIM, (MIKE11, NAM, HEC-HMS...). Hydraulic study of river systems uses 1 and 2-dimensional models (MIKE 11-ST, HEC-RAS, SOBEK ...). River basin and river estuary research often uses 2 and 3-dimensional models as these models will simulate the field dynamics of the study area (MIKEFLOOD, DELFT3D, EFDC...). The selection and comparison of computational models should be carried out to produce accurate calibration results for the study area. The mathematical model helps us overlay information in the study area continuously over time to simulate the state and predict the change of natural conditions.

MIKEFLOOD is one of modules in MIKE package. It is integrated from 1-D flood modeling (MIKE11) and 2-D surface flow (MIKE21) to calculate unsteady flow hydraulics. The simulation results from this module have been verified in many basins in the world. In this study, MIKEFLOOD combined with Geographic Information System (GIS) is selected.

To compute the hydraulic boundaries, MIKE11HD has applied the NAM model. NAM is the abbreviation of the Danish (Nedbor - Astromnings - Model) which means precipitation - runoff flow. This model was originally developed by the Department of Hydrodynamics and Water Resources at the

Technical University of Denmark (Nielsen and Hansen, (1973)). This is a conceptual model, describing the physical characteristics of the basin, on the basis of which it calculates rainfall flows. NAM is the conceptual hydrological model with concentration parameter. Its parameter and variables present the mean values for the entire basin. Several model parameters can be evaluated from physical characteristics of the basin but the result indicators must be obtained by comparing the calculated flow and the measured actual flow (fig. 2).



Fig. 2. The structure of NAM rainfall-runoff model

Theoretical background of MIKE11HD model [5]

Hydraulic equation: MIKE11HD employs full dynamic description for solving equations for conserving kinetic energy and integrating vertically (Saint Venant equations).

$$\frac{\partial Q}{\partial x} + \frac{\partial A}{\partial t} = q$$
$$\frac{\partial Q}{\partial t} + \frac{\partial}{\partial x} \left(\alpha \frac{Q^2}{A} \right) + gA \frac{\partial h}{\partial x} + \frac{gQ|Q|}{C^2 AR} = 0$$

Where: Q is the flow rate; A is the area of the wet cross-section; q is lateral flow; h is the

water depth; C is the Chezy coefficient; R is the hydraulic radius; g is the gravitational acceleration.

The combined model MIKE21 is a dynamic modeling system that can be applied to estuaries, coastal areas, and rivers. The model includes the following modules: Flow module; Diffusion load module; Water quality module and ecology module; Modular transport of sludge, sand; Wave spectrum module.

Theoretical background of MIKE21 model [4, 6]

MIKE21HD module mainly calculates unsteady flow in relatively shallow water. It takes into account the effects of tides, winds, atmospheric pressure, density variations (temperature and salinity), waves, tangles, and wet points on river banks. The equation of flow in this module is as follows:

Flow simulation equation.

The equation for simulating the flow of 2-dimensional flow is integrated from the vertical 3-dimensional equation showing the material conservation processes (1) and the momentum (2), (3) of the flow as follows:

$$\frac{\partial \varsigma}{\partial t} + \frac{\partial p}{\partial x} + \frac{\partial q}{\partial y} = 0 \tag{1}$$

$$\frac{\partial p}{\partial t} + \frac{\partial}{\partial x} \left(\frac{p^2}{h} \right) + \frac{\partial}{\partial y} \left(\frac{pq}{h} \right) + gh \frac{\partial \zeta}{\partial x} + gp \frac{\sqrt{p^2 + q^2}}{C^2 h^2} - \frac{1}{\rho_w} \left[\frac{\partial}{\partial x} (h\tau_{xx}) + \frac{\partial}{\partial y} (h\tau_{xy}) \right]$$
(2)
$$-\Omega q - fVV_x + \frac{h}{\rho_w} \frac{\partial}{\partial x} (p_a) = 0 \frac{\partial q}{\partial t} + \frac{\partial}{\partial y} \left(\frac{q^2}{h} \right) + \frac{\partial}{\partial x} \left(\frac{pq}{h} \right) + gh \frac{\partial \zeta}{\partial y} + gp \frac{\sqrt{p^2 + q^2}}{C^2 h^2} - \frac{1}{\rho_w} \left[\frac{\partial}{\partial y} (h\tau_{yy}) + \frac{\partial}{\partial x} (h\tau_{xy}) \right]$$
(3)
$$-\Omega p - fVV_y + \frac{h}{\rho_w} \frac{\partial}{\partial y} (p_a) = 0$$

In the equation symbols are used as follows: h(x, y, t) - the water depth at point (x, y)relative to zero level; $\zeta(x, y, t)$ - water surface elevation relative to zero level; p,q,(x,y,t)the flux density in x and y directions $(m^3/s/m =$ (*uh*, *vh*); (*u*, *v*)- Average velocity in depth in the x and y directions; C(x, y, t) - Chezy $(m^{1/2}/s);$ coefficient ggravitational acceleration; f(V) - wind friction coefficient; $V, V_x, V_y(x, y, t)$ wind speed in x and y directions; $\Omega(x, y)$ - Coriolis (s⁻¹); $p_q(x, y, t)$ - atmospheric pressure (kg/m/s²); ρ_w - water density (kg/m^3) ; x, y- coordinates (m); t- time (s); $\tau_{xx}, \tau_{yy}, \tau_{yy}$ - shear stress components.

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The alternating direction implicit (ADI) approach and the double-precision (DS) algorithm for each directional matrix are used to solve the continuous-equation system and to preserve the flow of momentum.

Setting up MIKEFLOOD mathematical model for the study area

Input data.

Evaporation: Data at Dong Hoi meteorological station is used. Only Piche evaporation is available, therefore calculation of ETo (by Penman formula) is applied.



Fig. 3. Connecting NAM, MIKE1HD and MIKE21FM (MIKEFLOOD) models

Rainfall: Rainfall data is available at Kien Giang station (November 1999, November 2000, November 2015, Obtober 2016). (Source: General Department of Irrigation).

Discharge: The observed hourly flood level in 2016 is available at Kien Giang Station. Based on $Q \sim H$ curve at Kien Giang station, the hourly discharge in Obtober 2016 is obtained. Flood event in November 2015 is used for verification. After calibration and verification of parameters in NAM model, rainfall data in Dong Hoi station are used to simulate flood flows in November 1999 and November 2000, and these results are used as inputs for the MIKE11 HD model.

Water level: Observed data at Dong Hoi Station (November 1999, November 2000).

Topological data: DEM in 30 m resolution is gridded in 10 m for MIKE21, and it is also used to create cross sections for MIKE11 model (fig. 3). DEM is created from topographical map at scale of 1:10,000. (Source: Institute of Geography).

Remote Sensing and Geographic Information System (GIS)

Method of interpolation of waterways [8, 9]

This article uses a combination of NDWI and the ratio method used by the Green/NIR/Green image channel and the medium/Green/MIR because the ratio of Green/NIR is necessary to extract the soil from the plant and the Green/MIR ratio is necessary to separate vegetation and unused land.

In which, the NDWI (Normalized Differenced Water Index) is calculated according to Gao's formula [8] as follows:

$$NDWI = \frac{B4 - B5}{B4 + B5} \tag{1}$$

To separate the waterline on Landsat images, the authors used the Gonçalves formula [10]:

$$Shoreline = \frac{B2}{B4} \times \frac{B2}{B5}$$
(2)

In which: *B*2: Is the Green channel on the Landsat image (Band 2 for Landsat 5 TM, Band 3 for Landsat 8 OLI); *B*4: Red channel on Landsat image (Band 3 on Landsat 5 TM image, Channel 4 on Landsat 8 OLI image); *B*5: Infrared channel on Landsat image (Band 4 on Landsat 5 TM image, Channel 5 on Landsat 8 OLI image).

Resolving clouds on Landsat image solving

The Landsat image was studied by the Canadian Remote Sensing Center (Haze Optimized Transform) [10]. The HOT method is based on the fact that the bands in the wavelength range seen are highly correlated to the subjects in the sky shooting but vary in the degree of cloudiness of the gray value brother.

The unclouded areas of the scene are visually identified and used to identify the "inner row" of relations between bands with a variety of surface cover factors. The HOT algorithm is measured by the deviation of the orthogonal in the spectral domain of the visible wave, *B*1, *B*3 is the grayscale values of bands 1 (Blue) and 3 (Red). HOT is calculated by the following formula:

$HOT = B1sin\theta - B3cos\theta$

The angle θ is determined from the highly correlated pixel patterns (between *B*1 and *B*3) in the cloudless region of the scene. HOT values are calculated for each pixel then considered to be the spatial distribution of the cloud. A HOT photo mask will be used to eradicate haze at some points, even in some of the undetected photos with the naked eye. The adjustable gray level for each pixel is different between the lower limit of the histogram involved and the histogram of the pixels in the cloudless region. Excluding cloud cover can be determined not depending on the level of the original cloud. Although every pixel affected by haze is highlighted, only some can be calibrated.

RESULTS AND CONCLUSION



Fig. 4. Separation of water body by *NDWI* index from Kien Giang - Quang Binh river section

After implementing the above method, the shoreline will be made in raster format, then the data will be transferred to GIS software to semi-automatic and the shoreline map will be completed. Applied to the Kien Giang basin the article uses the image of Landsat 5 TM with 2 scenes taken on November 11, 1999 and November 5, 2000 and Landsat 8 OLI taken on November 15, 2015 downloaded from the website: http://glovis.usgs.gov/ [9] (fig. 4).

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The result from MIKEFLOOD shows that the flow rate calculated from the MIKE11-NAM model is in quite good agreement with observed discharge (fig. 5). NASH coefficient is very high (98% - table 1). The parameter set (table 2) therefore is used to simulate runoff for subsequent years which is used as boundary input for hydraulic modeling for routing flow in the river.



Fig. 5. Calibration and verification results of dischage at Kien Giang station

Table 1. Flood path selection for calibration and verification cases

	Time	NASH
Calibration	from October 1 to November 23, 2016	0.98
	from October 1 to November 23, 2015	0.89
Verification	from November 1 to November 11, 1999	0.87
	from November 1 to November 15, 2000	0.78

Table 2. NAM model setting,	flood event	October 1	to November	23, 2016 at I	Kien Giang sta	tion
				,	0	

Umax	Lmax	CQOF	CKIF	CK1,2	TOF	TIF
12.6	101	0.35	550.1	11.8	0.36	0.501

After calibration and verification, simulations of the floods on November 11, 1999 and November 1, 2000 are carried out to compare with Landsat 5 satellite image. The simulation result from the MIKEFLOOD and the processed image show a fairly similar inundation. The similarity is quite high in some communes such as Hong Thuy, Loc Thuy, Kien Giang towns (fig. 6, 7, 8, 9, 10 and 11). An Thuy commune has the largest flooded area (more than 2,000 ha), and Hong Thuy commune has a deep flooding level (deeper than 3 m). This is a low lying area so drainage is not easy. The area near My Trung dam is shrinking so floods are frequent (table 3, 4 and 5).



Fig. 6. Result of flooded area from the MIKEFLOOD model and Landsat 5 satellite imagery of flood November 11, 1999



Fig. 7. Result of flooded area from the MIKEFLOOD model and Landsat 5 satellite imagery of flood November 5,2000

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Fig. 8. Result of flooded area from the MIKEFLOOD model and Landsat 8 satellite imagery of flood November 15,2015



Fig. 9. Comparison of flooded area from the MIKEFLOOD model and satellite imagery of flood November 11, 1999



Fig. 10. Comparison of flooded area from the MIKEFLOOD model and satellite imagery of the flood on November 5, 2000



Fig. 11. Comparing the flooded area from the MIKEFLOOD model and the satellite imagery of the flood in November 15, 2015

Commune	-	Flood area from Satellite image									
	0.1 - 0.5	0.5 - 1	1 - 1.5	1.5 - 2	2 - 2.5	2.5 - 3	3 - 3.5	3.5 - 4	4 - 4.5	Total (ha)	Total (ha)
An Ninh	147.6	182.0	521.0	207.2						1057.8	1374.7
An Thuy	152.7	302.1	444.8	734.3	333.9	33.7	6.4	2.8	1.3	2012.0	2112.4
Gia Ninh	80.6	124.9	224.7	389.2	29.6	0.2				849.3	928.0
Hong Thuy	45.0	55.8	107.3	428.0	525.7	58.0	22.3	19.0	6.6	1267.7	1315.8
Hien Ninh	110.3	151.5	144.3	23.8	0.7					430.6	534.2
Hoa Thuy	42.7	90.2	128.1	420.4	292.9	0.1				974.5	1246.0
Loc Thuy	41.5	80.2	226.9	59.8	5.2	0.3				414.0	415.9
Lien Thuy	131.9	139.8	126.1	44.4	12.7	6.3	1.5			462.8	665.5
Mai Thuy	102.1	76.5	72.6	59.5	37.5	7.7	2.3	0.9	0.8	359.9	661.6
My Thuy	73.3	51.3	32.8	32.5	28.3	18.5	9.0	1.6	0.7	247.9	790.8
Phu Thuy	162.1	307.3	264.1	118.2	21.7	1.7				875.1	1271.3
Phong Thuy	35.5	88.2	367.6	394.5	64.6	1.9				952.3	939.5
Son Thuy	133.5	107.1	111.3	222.0	34.2					608.1	871.9
Tan Ninh	204.0	420.2	364.0	96.7						1085.0	1158.6
Thanh Thuy	57.2	73.1	141.5	268.1	32.9					572.9	595.6
Truong Thuy	25.4	18.6	15.3	14.0	10.8	9.7	8.5	7.8	4.0	114.0	43.5
Kien Giang	77.6	78.5	82.2	56.1	16.8	2.3	0.0			313.6	364.7
Van Thuy	11.6	9.3	6.7	6.1	6.8	5.0	1.8	0.2		47.4	74.7
Van Ninh	65.0	75.3	204.6	471.0	24.8					840.7	1073.7
Vo Ninh	210.1	164.9	58.0	24.0	3.6	2.1	1.4	0.7	0.2	465.1	794.9
XuanThuy	81.2	128.3	86.7	36.9	18.5	9.8	2.7	0.2		364.3	438.9

Table 3. Flooded areas (ha) and flood levels (m) during flood on November 11, 1999

Table 4. Flooded areas (ha) and flood levels (m) during flood on November 5, 2000

Commune		Flood area from Satellite image									
	0.1 - 0.5	0.5 - 1	1 - 1.5	1.5 - 2	2 - 2.5	2.5 - 3	3 - 3.5	3.5 - 4	4 - 4.5	Total (ha)	Total (ha)
An Ninh	163.4	541.1	160.7							865.3	698.5
An Thuy	349.4	399.0	623.7	188.4	13.6	3.6	1.5	1.0	0.9	1581.2	1940.7
Gia Ninh	126.7	191.3	368.4	21.7	0.0					708.2	719.2
Hong Thuy	53.0	107.6	461.8	501.2	52.3	22.5	17.8	3.6	2.0	1223.5	1145.1
Hien Ninh	86.7	63.3	4.7	0.2						154.9	230.0
Hoa Thuy	93.1	140.5	484.4	211.0						929.0	932.8
Loc Thuy	84.1	230.1	60.2	5.6	0.3					380.2	337.9
Lien Thuy	91.7	109.4	37.2	12.5	6.5	1.8				259.0	367.4
Mai Thuy	63.8	40.4	12.6	3.1	1.6	0.1				121.6	267.0
My Thuy	29.5	29.3	21.8	19.1	11.6	4.1	0.6	0.0		115.9	252.8
Phu Thuy	155.0	104.7	21.4	1.8						282.9	812.7
Phong Thuy	88.3	376.4	387.0	63.5	1.8					916.9	730.2
Son Thuy	63.3	104.6	212.0	18.6						398.6	576.1
Tan Ninh	328.0	351.1	62.7							741.7	531.6
Thanh Thuy	81.8	147.2	265.9	30.7						525.6	438.5
Truong Thuy	15.5	12.2	6.7	3.0	0.6	0.0				38.0	41.9
Kien Giang	65.1	43.0	23.5	13.3	3.2	0.2				148.2	248.2
Van Thuy	6.3	6.6	6.5	2.2	0.5	0.0				22.2	55.9
Van Ninh	83.5	245.9	444.5	10.8						784.7	774.8
Vo Ninh	163.4	541.1	160.7							865.3	698.5
Xuan Thuy	68.2	33.8	9.2	3.6						114.8	254.5

Commune		Flood area from Satellite image									
	0.1 - 0.5	0.5 - 1	1 - 1.5	1.5 - 2	2 - 2.5	2.5 - 3	3 - 3.5	3.5 - 4	4 - 4.5	Total (ha)	Total (ha)
An Ninh	374.0	333.6	0.7							708.3	698.5
An Thuy	333.7	503.2	455.6	34.7	5.9	2.0	1.2	1.1	0.6	1337.9	1940.7
Gia Ninh	133.5	394.4	61.4	0.7						589.9	719.2
Hong Thuy	77.7	288.8	686.0	76.1	25.9	20.5	6.3	2.0	1.6	1185.9	1145.1
Hien Ninh	74.6	22.0	2.2	0.2						99.0	230.0
Hoa Thuy	103.5	340.6	405.7	1.0						850.9	932.8
Loc Thuy	181.1	141.5	17.3	0.9	0.0					340.9	337.9
Lien Thuy	102.5	56.6	11.4	5.9	4.9	2.4				183.7	367.4
Mai Thuy	67.3	52.0	27.7	4.3	1.9	0.6	0.1			154.1	267.0
My Thuy	39.1	32.6	24.0	18.1	15.6	5.2	1.0	0.0		135.5	252.8
Phu Thuy	183.9	78.7	11.6	0.6						274.8	812.7
Phong Thuy	220.4	513.8	126.2	8.9	0.5					869.7	730.2
Son Thuy	64.0	211.4	61.6							337.0	576.1
Tan Ninh	305.1	172.4	0.4							477.9	531.6
Thanh Thuy	96.3	262.3	82.4							440.9	438.5
Truong Thuy	17.8	13.1	9.6	5.7	2.6	1.7	0.1			50.7	41.9
Kien Giang	69.3	60.1	19.7	13.0	3.9	0.3				166.3	248.2
VanThuy	6.8	6.3	7.6	4.2	1.5	0.2				26.7	55.9
Van Ninh	156.7	476.5	96.9							730.1	774.8
Vo Ninh	87.8	16.1	4.7	3.7	2.2	1.6	0.8	0.3		117.1	381.3
Xuan Thuy	132.2	59.5	18.3	12.5	9.1	2.1	0.3			234.0	143.4

Table 5. Flooded areas (ha) and flood levels (m) during flood on November 15, 2015

The simulation results from the model and remote sensing images can be converted to vector format (GIS shapefile) to calculate the flooding area and flooding level at each commune. The combination of these two methods shows that it is possible to accurately evaluate flooding in the study area. The mathematical model can help to predict flooding area and flooding level in detail while the remote sensing image shows the actual flooding visually.

Every year, Kien Giang river basin is affected by three to four storms of high intensity. As a result, the basin is regularly flooded, affecting the livelihood of local people.

The study has applied the MIKEFLOOD combined with Geographic Information System to calculate and simulate flood inundation in the downstream area of Kien Giang river. At the same time, remote sensing technology is used to analyse the situation at the time of flood in order to compare, and verify the calculation results from the mathematical model. The results from the two methods are quite similar. This result will be the scientific basis for making plan of flood prevention as well as of socio-economic development strategies for Quang Binh province.

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