

Assessment of geomorphic processes and active tectonics in Con Voi mountain range area (Northern Vietnam) using the hypsometric curve analysis method

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ABSTRACT

The main objective of this study is to assess geomorphic processes and active tectonics in the Day Nui Con Voi (DNCV) area of Vietnam. For this purpose, a spatial database was collected and constructed, including DEM (Digital Elevation Model) and a geological map. The hypsometric curve (HC) analysis method and its statistical moments were adopted to use for the assessment. These methods have been widely used for the assessment of geomorphic processes and active tectonics in many areas in the world showing promising results. A total of 44 sub-basins of the Red River and the Chay river were analyzed. The result shows that 3 curve-types such as "straight- shape", "S-shape", and concave were found; with the concave curve being the dominant and widely distributed in the northeast side and in the south of the southwestern side of the study area. The hypsometric integral (HI) values are rather small with the largest value is 0.37 and the smallest one is 0.128. Other statistical moments of the hypsometric curve, i.e. skew (SK), kurtosis (KUR), and the density function (density skew - DSK and density kurtosis-DKUR) show great values, which increased in the south direction of the area study. Accordingly, recent active tectonics (uplift-lower) in the study area is generally weak; however, they are also not completely homogeneous and can be distinguished by different levels. The southwestern side is being lifted higher than the northeastern side. The northern part is being lifted larger than the southern part. In the region, the uplift activities were increased gradually in the Pliocene-Quaternary and could have stopped at certain time in the past. The current geomorphic processes are mainly headward erosion in the upstream.

Keywords: Geomorphic index; Hypsometric curve; Statistical moments; Active tectonics; Red River fault; Day Nui Con Voi.

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1. Introduction

The Red River shear zone (RRSZ) extends

over a length of 1000 km from Tibet to the East Vietnam Sea. Along the shear zone, four narrow massifs of high-grade metamorphic complexes, the Day Nui Con Voi in Vietnam,

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Ailao Shan, Diancang Shan and Xuelong Shan in Yunnan, China are considered as the "axes" of the RRSZ - important geological boundaries in Asia. The Day Nui Con Voi range is in the southeasternmost part of this shear zone (Figure 1). This area has been received attentions of many geoscientists and seen as a key to understand the geodynamics of the RRSZ (Leloup et al., 1995; 2001; Le et al., 2004). The achieved results have contributed to the explanation and clarification of many issues in geology, tectonics and geomorphology. However, some points are not consistent and disputed (e.g. Tran et al., 1999; 2002; Le, 2003; Le et al., 2001; Phan et al., 2004; Wang et al., 2000; Leloup et al., 2001). Studies of tectonics in this area have not paid much attention to the role and significance of geomorphology; especially, the lack of quantitative analyses of landscapes using various geomorphic indices.

Geomorphic indices have been found to be useful in identifying areas experiencing tectonic activity because they facilitate rapid evaluation of large areas (Strahler, 1952; Bull and McFadden, 1977; Keller and Pinter, 2002; Joshi et al., 2013). Furthermore, active faults and growing folds commonly have topography that is useful in identifying different geomorphic or structural segments along the fault and estimating the most active segments (Azor et al., 2002; Font et al., 2010; Joshi et al., 2013). Segments along a morphostructure may be outlined and identified to determine the relative intensity of tectonic activity along a fault by utilizing a detailed study of drainage anomalies coupled with geomorphic indices (Azor et al., 2002; Keller and Pinter, 2002; Joshi et al., 2013). Moreover, with the current development of GIS, the calculation of geomorphic indices has become easier (Troiani and Della Seta, 2008; Pérez-Peña et al., 2009; Joshi et al., 2013). So, the geomorphic indices have been widely used in geomorphology and active tectonics (e.g., see in the above references).

In Vietnam, despite some initial geomorphic indices also to be used quite successfully in several studies such as Nguyen et al.

1999; Phung, 2011; Phan, 2014; Nguyen, 2015. However, most of the calculations in these studies were manually carried out based on topographic maps and satellite images; so the results often depend on the ability to estimate, sight and experience of experts who conducted these studies. Therefore, the analysis and assessment of geomorphic indices have not been shown clearly roles, the significances, and its relationship to the geomorphological processes and active tectonics.

In this paper, we present quantitative analyses and assessments of the hypsometric curve (HC) and its statistical moments in relationship between geomorphic processes and active tectonics in the DNCV area. The HC index is one of the geomorphic indices that has been considered as a powerful tool for quantifying the topographic features and differentiate zones deformed by active tectonics (Keller and Pinter, 2002; Chel et al., 2003; Pérez-Peña et al., 2009; Pedrera et al., 2009; Mahmood and Gloaguen, 2012). However, in Vietnam, this is the first time the method is adopted for the assessment of the active tectonics in the Lo River fault zone and the Tam Dao area (Ngo et al., 2016), but statistical moments of the hypsometric curve has not been analyzed and assessed.

2. Tectonic, geologic, and geomorphic settings

The Day Nui Con Voi (DNCV) mountain is less than 10 km wide and more than 250 km long, extending from Lao Cai to Viet Tri, and appearing as an elongated NW-trending core of metamorphic rocks (Tran et al., 1998) (Figure 2). The altitude of the mountain is peaked at Nui Lai of 1450 m, then descending to the northwest and southeast. This mountain is characterized by three main strips, with the NW-SE direction and separated by the parallel lines with the Red River. The topography in this area is asymmetry: slope of the northeastern side is smaller than the southwest side; on the northeastern side have some narrow strips extending along the main mountain; the southwest side is divided into individual peaks. The center strip of the DNCV is uplifted (500-1000 m) compared with the two sides (<500 m) (Le et al., 2004).

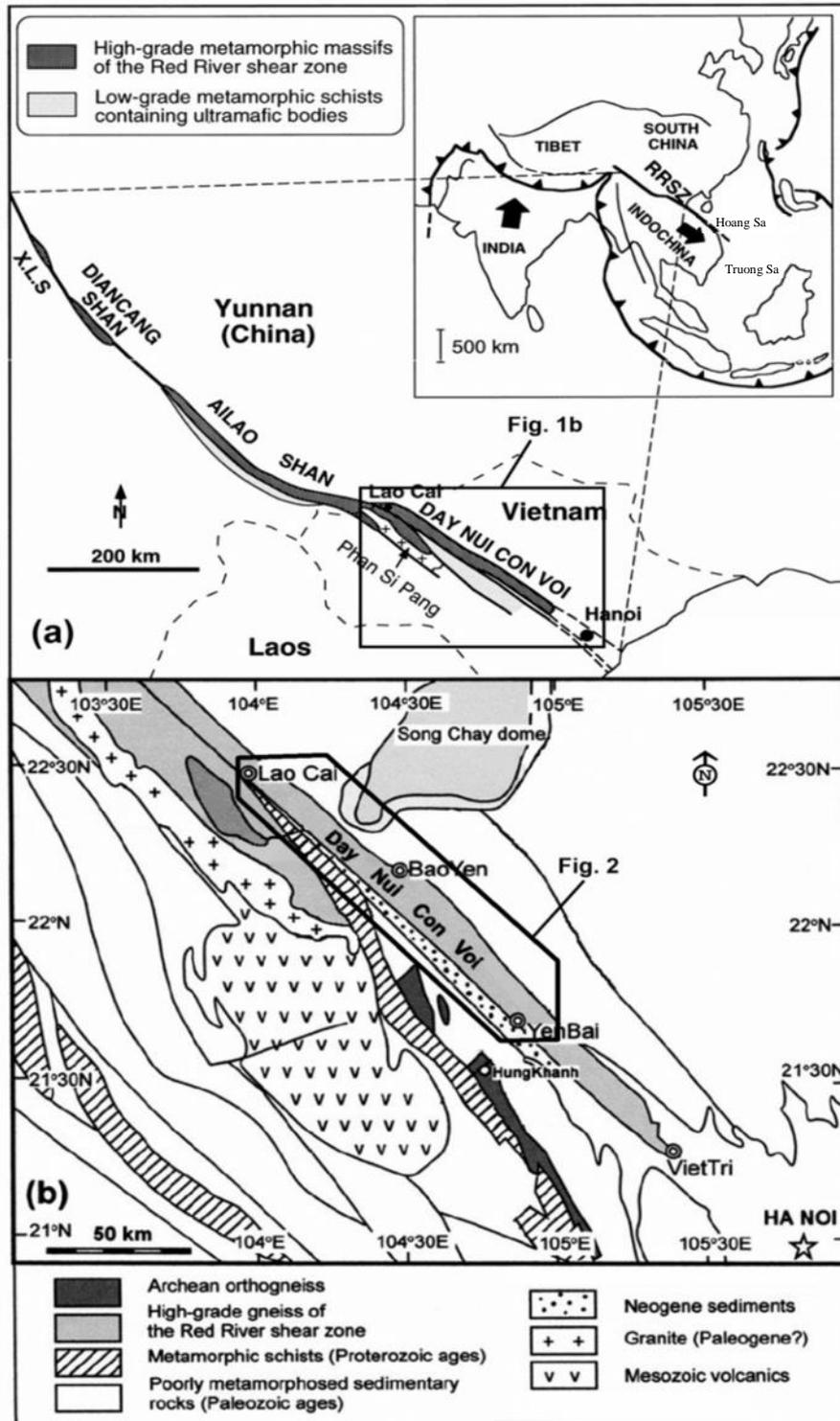


Figure 1. (a). The Red River shear zone in Asia, (b) geological sketch map around the Day Nui Con Voi (Modified after Tran et al., 1998; 2003)

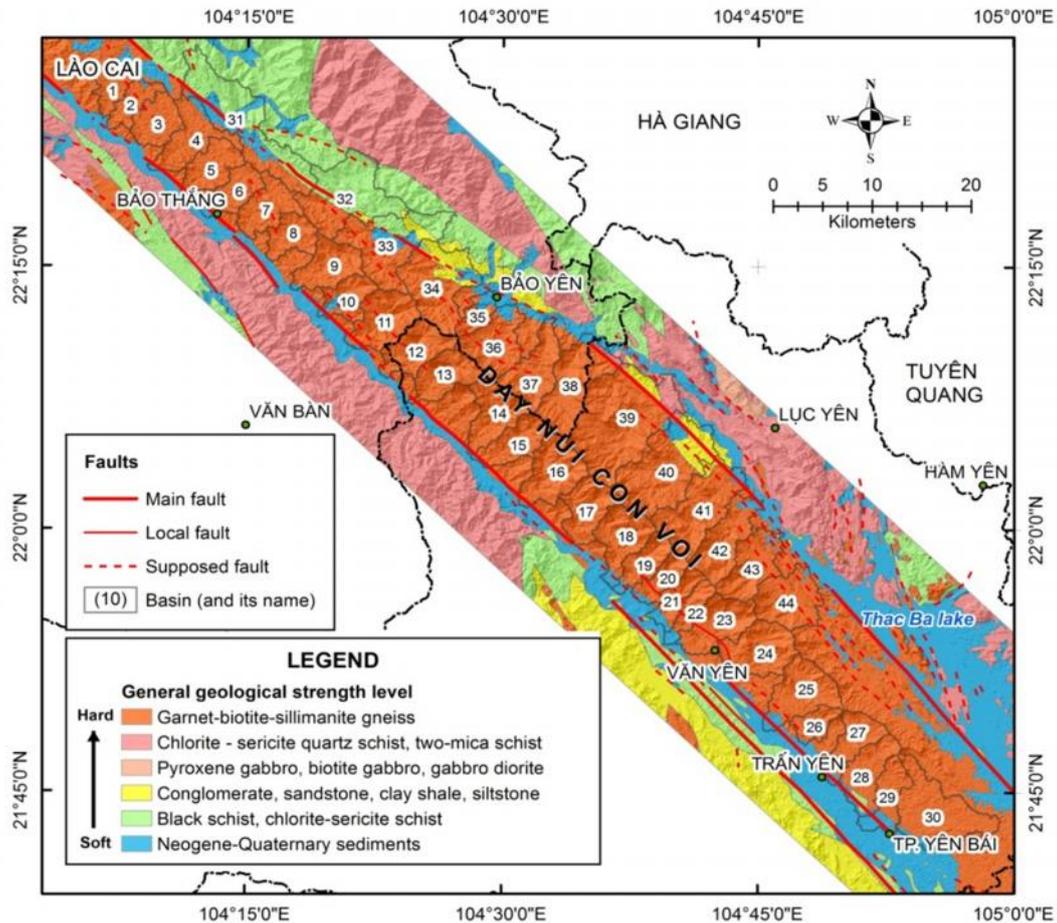


Figure 2. Geological strength level map in the Day Nui Co Voi and surrounding area

As for the Ailao Shan, the DNCV is a narrow high-grade metamorphic rocks and are mapped as Proterozoic (Phan et al., 1994; 2012). It is composed chiefly of garnet-biotite-sillimanite gneiss and garnet-biotite gneiss, and minor two-mica schists with garnet. The DNCV also includes amphibolite layers, migmatites, mylonite bands and small lenses of marble. This rock assemblage suggests that the DNCV was formed with severe deformation and deep metamorphism of sedimentary rocks (Tran et al., 1998; Phan et al., 1994; 2012). The rocks within the DNCV are strongly foliated. The foliation, which is marked by the preferred orientation of planar minerals (biotite and amphibole) and by flattened quartz or feldspar ribbons, commonly strikes parallel to the local trend of

the gneiss core and dips steeply (~70°) to the northeast. The lineation is deduced by elongated quartz and feldspar ribbons, long tails of feldspar porphyroblasts, stretched leucocratic veins and preferred orientations of sillimanite crystal shapes all locally plunge to the northwest in a range of 5-20° (Tran et al., 1998). A mylonite band about 200-500 m wide is well exposed in the center of the northeastern flank of the shear zone. Foliation and lineation within the mylonite band are parallel to those of the host gneisses. Numerous kinematic indicators suggest a left-lateral shear movement of this mylonite band (Phan et al., 1995; Tran et al., 1998). The foliation of gneisses is then cut by two sets of steep conjugate faults, N10°E striking dextral and more numerous N110°E striking sinistral,

indicating N60°E shortening. It shows that a successive deformation with ENE shortening (Tran et al., 1998).

From Vietnam-China border, at Lao Cai, the Red River valley fault splays into two roughly parallel strands, the Chay River and Red River faults, which bound the DNCV to the north and south, respectively. Currently, both fault-strands appear to slip mostly right-lateral slip, with variable components of normal slip (Allen et al., 1984; Phan et al., 1994, 2004, 2012). Narrow straight ‘grabens’, which are traced along the Red River and Chay River faults, are filled with Late Miocene sediments containing abundant pebbles of gneisses and mylonites, being interpreted as a synorogenic formation resulting from the reversal of fault movements from left-lateral to right-lateral about 5 m.y. ago (Leloup et al., 1994). On the SW and NE sides of the DNCV also exist some small faults run nearly parallel with the Red River and Chay River faults, respectively (Le et al., 2004).

3. Data and methods

To determine the hypsometric curve and its statistical moments for the study area, we used Digital Elevation Model (DEM) with 30 m resolution which is provided by the United States Geological Survey (USGS). The DEM is analyzed by ArcGIS software; it is useful tools to ensure accuracy, quick and less

expensive in the calculation of morphology parameters. The calculation in this study is carried out automatically using the extension tools of ArcGIS 10.1 software (Pérez-Peña et al., 2009). Geological map of the study area was constructed using the digital Geological and Mineral Resources maps at the scale of 1:200,000 (The Department of Geology and Minerals of Vietnam). We used the active faults from the Phan et al. (2004, 2012), Ngo et al. (2006, 2011), and Le et al. (2004).

3.1. Hypsometric curve and hypsometric integral

The hypsometric curve describes the distribution of elevations across an area of land with different scales from one drainage basin to the entire planet. The curve is created by plotting the proportion of total basin height ($h/H =$ relative height) against the proportion of total basin area ($a/A =$ relative area) (Strahler, 1952; Keller and Pinter, 2002) (Figure 3). The shape of the hypsometric curve is related with the stage of geomorphic development of the basin. Convex hypsometric curves are typical of a youthful stage; S-shaped curves are related to a maturity stage, and concave curves are indicative of a peneplain stage (Strahler, 1957; Gardner et al., 1990; Delcaillau et al., 1998; Keller and Pinter., 2002; Pérez-Peña et al., 2009) (Figure 3).

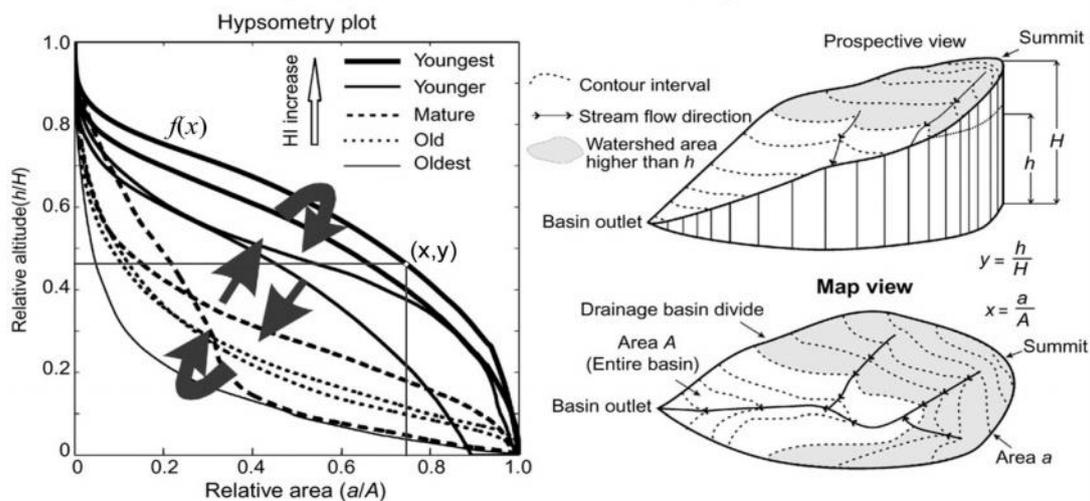


Figure 3. Basic hypsometric curves and its geomorphological development cycles (Modified after Strahler, 1952; Pérez-Peña et al., 2009; Mahmood and Gloaguen, 2012)

A simple way to characterize the shape of the hypsometric curve for a given drainage basin is to calculate its hypsometric integral (HI). The integral is defined as the area under the hypsometric curve and can be calculated (Keller and Pinter, 2002):

$$HI = (H_{\text{mean}} - H_{\text{min}}) / (H_{\text{max}} - H_{\text{min}}) \quad (1)$$

where HI is hypsometric integral, H_{max} is maximum elevation, H_{min} is minimum elevation, and H_{mean} is mean elevation.

The parameters in the formula (1) can be identified by analyzing the DEM with the GIS software. The HI index has been used, as well as the hypsometric curve, to infer the stage of development of a basin. The values of the HI always vary from 0 to 1. Values near 1 indicate a state of youth and are typical of convex curve. However, in the mature stage of the basin, it has a lot of S-shape and concave shape but the HI values often similar. Meanwhile, to distinguish or assessment correlate between the basins, we often base on the statistical indices are given below.

3.2. Statistical moments of the hypsometric curve

In addition to analyzing hypsometric integral (HI) index, we also calculate and analyze other statistic moments of hypsometric curve (HC): skewness of the hypsometric curve (hypsometric skewness, SK), kurtosis of the hypsometric curve (hypsometric kurtosis, KUR), skewness of the hypsometric density function (density skewness, DSK), and kurtosis of the hypsometric density function (density skewness, DKUR).

Harlin (1978) developed a technique that treated the hypsometric curve as a cumulative probability distribution and used its statistic moments to describe it quantitatively. It consists of the hypsometric curve by a continuous polynomial function with the form (Harlin, 1978) (Figure 3).

$$f(x) = a_0 + a_1x + a_2x^2 + \dots + a_nx^n \quad (2)$$

and HI can be defined:

$$HI = \iint_R dx dy \quad (3)$$

where R is the region under the hypsometric curve, x is relative area, and y is relative height.

Skewness of the hypsometric curve is defined by:

$$SK = \mu_3 / (\mu_2^{1/2})^3 \quad (4)$$

where μ_3 and μ_2 are the third-order and second-order moment about x ,

$$\mu_3 = \frac{1}{HI} \iint (x - \mu_1)^3 dx dy \quad (5)$$

$$\mu_2 = \frac{1}{HI} \iint (x - \mu_1)^2 dx dy \quad (6)$$

where μ_1 is the first-order moment or x mean or x centroid,

$$\mu_1 = \iint_R x dy dx \quad (7)$$

Kurtosis of the hypsometric curve is defined by:

$$KUR = \frac{\mu_4}{(\mu_2^{1/2})^4} \quad (8)$$

where μ_4 is fourth-order moment about x ,

$$\mu_4 = \frac{1}{HI} \iint (x - \mu_1)^4 dx \quad (9)$$

Density skewness (DSK) and density kurtosis (DKUR) are defined similarly except that now y is the first derivative of the hypsometric curve, i.e., the density function of the hypsometric curve (replacing y with y'). These definitions are chosen so that they are consistent with Harlin's original work (Harlin, 1978).

In statistics, skewness and kurtosis describe the shape of a distribution relative to the normal distribution and are dimensionless. Skewness characterizes the degree of asymmetry of a distribution around its mean. A positive value of skewness ($SK > 0$) signifies a distribution with an asymmetric tail extending out toward a more positive x (skewed to the right); a negative value ($SK < 0$) signifies a distribution whose tail extends out toward a more negative x (skewed to the left); and the skew is zero ($SK = 0$), when the variable distribution is symmetrical. Kurtosis measures the relative peakedness or flatness of a distribution, relative to a normal distribution. Larger kurtosis ($KUR > 3$) indicates a "sharper"

peak than normal distribution (the same Luo, 2000 and Pérez-Peña et al., 2009, under the definition used in this paper, the kurtosis of a normal distribution is 3); smaller kurtosis indicates "flatter" peak than normal distribution.

These statistics are applied to the distribution function of the hypsometric curve order to explain the erosion and slope basins and has been tested by Harlin., (1978); Luo., (1998, 2000); Pérez-Peña et al., (2009). Accordingly, the hypsometric skewness represents the amount of headward erosion in the upper reach of a basin (Figure 4); density skewness indicates slope change; a large value of kurtosis signifies erosion on both upper and lower reaches of a basin, and density kurtosis delineates midbasin slope.

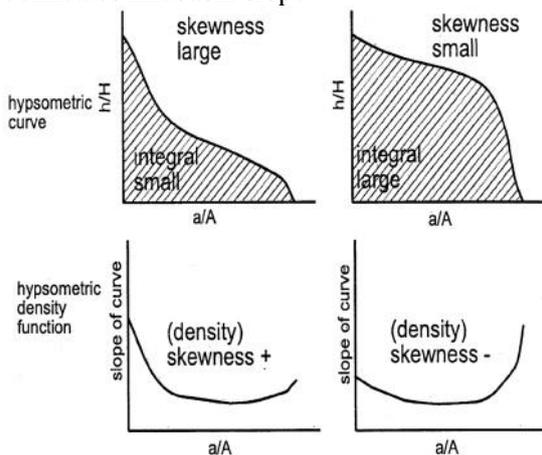


Figure 4. Schematic diagram showing the relationship between the shape of the hypsometric curve and its integral, skewness, and density skewness (Luo, 2000)

These statistical moments can be used to describe and characterize the shape of the hypsometric curve and, hence, to quantify changes in the morphology of the drainage basins. In many cases, these parameters are very useful for the hypsometric analysis, especially in basins with similar hypsometric integrals but different shapes (Pérez-Peña et al., 2009).

4. Results

In the DNCV area, the hypsometric curve analysis method and its statistical moments are used for assessment at 44 sub-basins of the

Red river and the Chay river. In which, 30 sub-basins are located in the Red River (from the basin 1 to 30) and 14 sub-basins are located in the Chay River (from the basin 31 to 44) (Figure 5). The results are showed on Table 1, Figures 5 and 6.

In the study area, the hypsometric curve can be grouped into 3 curves: "straight-shape", "S- shape", and concave curves (Figs. 6a, 6b and 6c,d, respectively) and no convex curve. Accordingly, concave curve has the largest proportion (26/44 basins), followed by the S-shape (10/44 basins) and final are straight-shape (8/44 basins). Consistent with them, the HI indices are also very small, the largest value is the basin No.13 (HI = 0.37) and the smallest is the basin No.28 (HI = 0.128). In which, the basins with "straight-shape" have the HI values are greater than 0.3; the "S-shape" have HI values are greater than 0.25 and the concave curves with largest HI value is 0.28 (Table 1).

The results shown in Table 1 show that the skew values are from 0.45 to 1.3 and these values do not change much in the basins with straight-shape of the hypsometric curve (the skew values range from 0.55 to 0.83) and the "S-shape" of the hypsometric curve ($0.45 < SK < 0.64$). In contrary, the skew values have considerable variability in the basins with concave shape of hypsometric curve (the skew values range from 0.46 to 1.3). In the basins with straight-shape and s-shape of hypsometric curve, the density skew values range from 0.33 to 0.96, and the basins have concave curve, this values range from ~ 0.78 to 1.58. The kurtosis values range from ~2.0 to 4.1; in there, the basins have the hypsometric curve with the "straight" and "S" shape, the kurtosis values are less than 3.0 (the kurtosis of a normal distribution is 3.0). The density kurtosis values range from 1.75 to 4.87. As the skew values, the density kurtosis values are not change much in the hypsometric curve basins with the "straight" and "S" shape, and quite change in the concave shape basins. The variation values of the main statistical moments indices in the DNCV are showed on Figure 7.

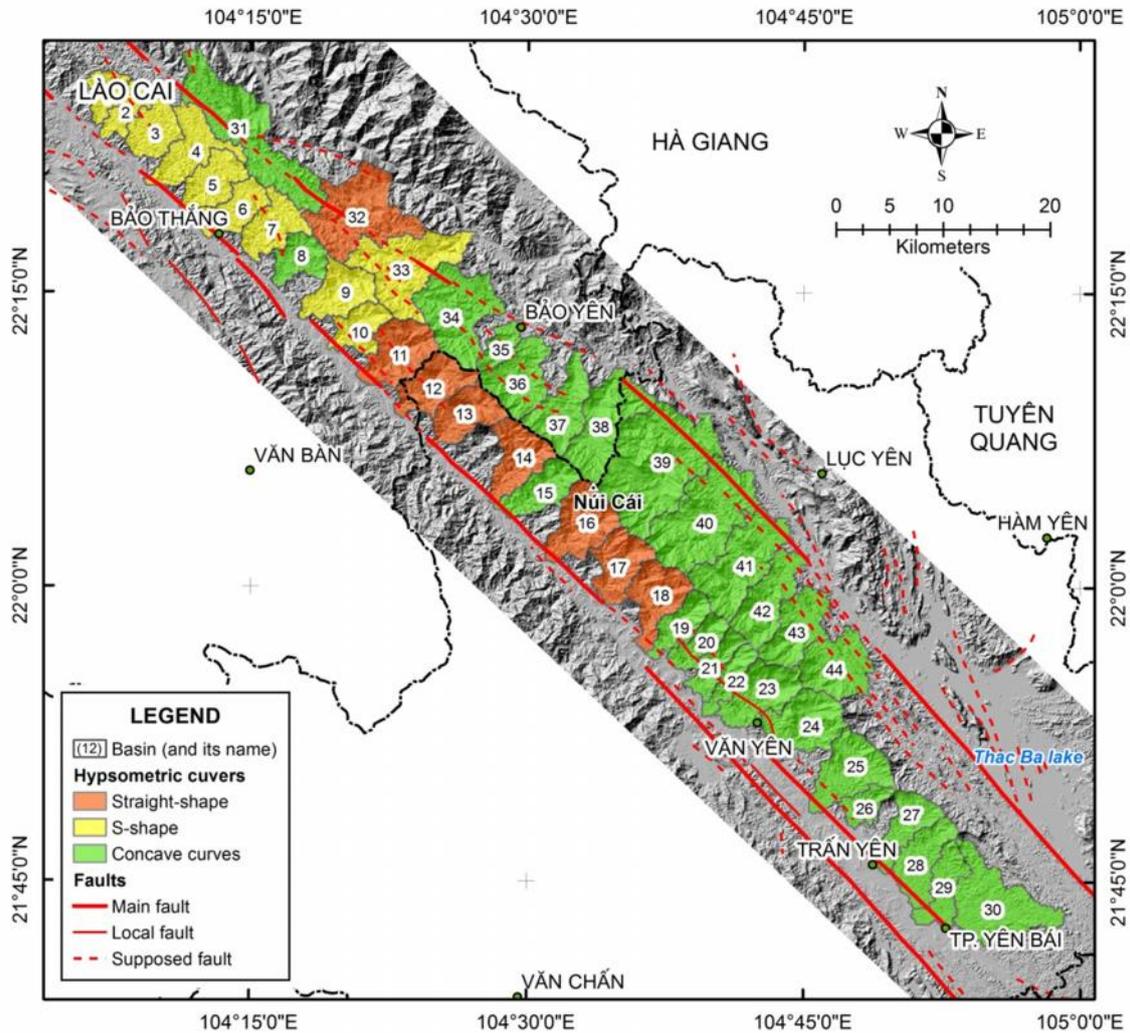


Figure 5. Schematic distribution of the hypsometric curve in the DNCV area

5. Discussion

The hypsometric curve and its statistical moments influenced by active tectonics, are also affected by geological and regional climatic characteristics (Moglen and Bras, 1995; Willgoose and Hancock, 1998; Huang and Niemann, 2006; Pedrera et al., 2009). Because the study area is located almost in the center of the DNCV with a narrow range, so the climate is basically not much different. According to the geological map (1:200,000) of the Department of Geology and Minerals of Vietnam, the DNCV area has identical geol-

ogy and is composed chiefly of high-grade metamorphic rocks (Figure 2). So, anomalies (if any) of geomorphic indices in this area are mainly a reflection of the recent tectonic activity.

Regarding to the difference of the number basins in the northeast side (14/44) and the southwest side (30/44) of the DNCV area, because in the southeastern part of this area has the Thác Ba hydropower dam, so the basins should flow directly into the lake having been changed base erosion level by the volume of water. Therefore, we did not use these basins in the calculations. On the other hand, due to

relief features of the DNCV with slopes in the southwestern side (in the Red River basin) is greater than the northeastern side (in the Chay River basin). Therefore, area of the basins in the southwestern side usually smaller than the northeastern side and opposite side, the number of basins in the northeastern side is

less than in the southwestern side. The steeper and higher of the southwestern side than the northeastern side reflected lift active of the DNCV in the southwestern side is higher than the northeastern side. This will be clarified by analyze the hypsometric curve and its statistical moments as below.

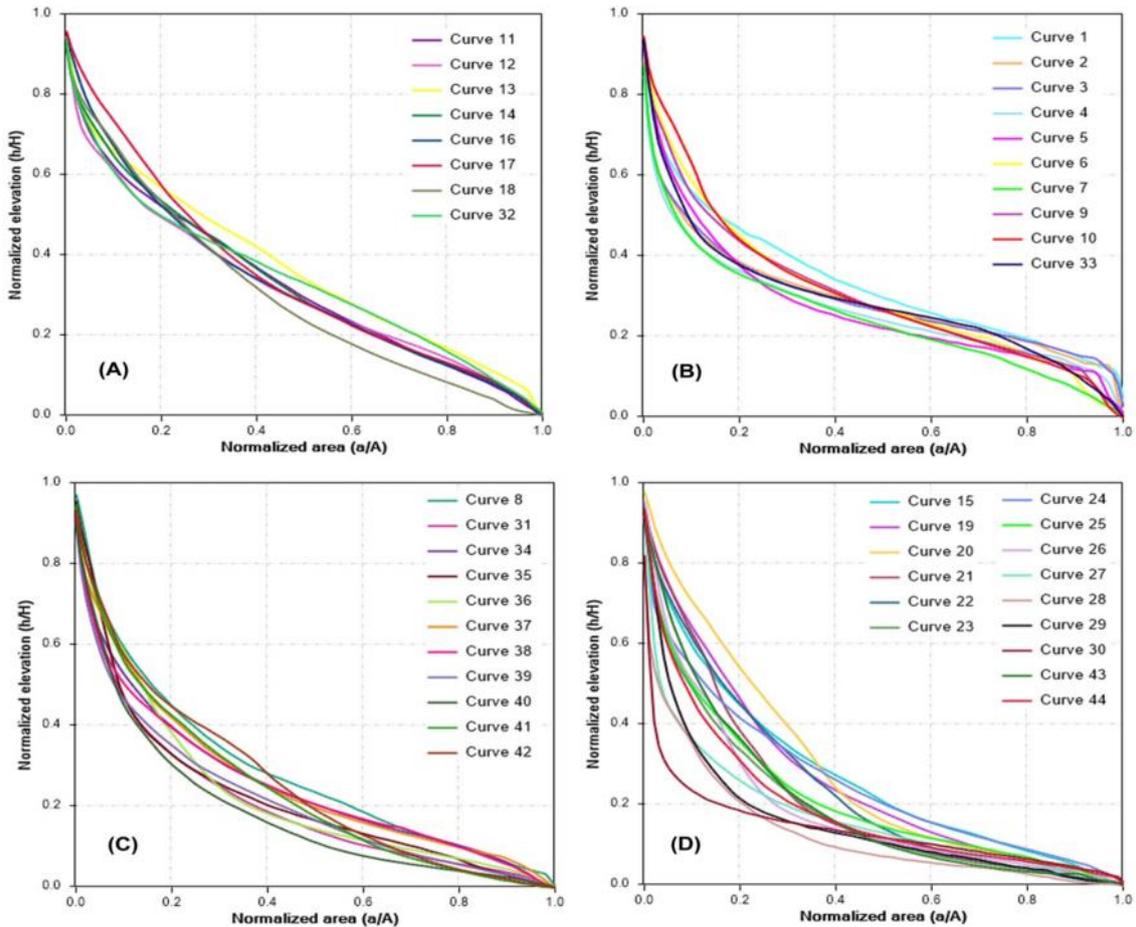


Figure 6. Hypsometric curves of the sub-basins in the DNCV area; (A) - “Straight-shape” group; (B)- “S-shape” group; (C) and (D)- concave curves

As the results presented above, in the study area, the hypsometric curve has revealed 3 curves such as "straight- shape", "S- shape", and concave curves, but no convex curve. In there, the hypsometric curve is almost concave curve (26/44 basins) and fit it, the HI values mainly small; maximum is 0.37 (Figure 5 and Table 1). Accordingly, the basin in this study area is mainly in the oldest stage,

meaning that the basin has reached the equilibrium in the longitudinal profiles of the river (or stream). In these basins, the dominant geomorphological processes usually are lateral erosion, vertical erosion (if any) also occurs in the upstream area. Another way, the active tectonics (uplift-lower) in these basins is basically weak. However, there still exists the hypsometric curve as "straight-

shape" and "S-shape" are distributed in some parts of the study area and focused mainly in the northern part to the center of the southwestern side of the DNCV. Whereas, in the northeastern side of the DNCV, the hypsometric curve mainly is concave curve (Figure 5 and 6a,b). Tectonic activity in the study area is not fully uniform. Accordingly, uplift active in the southwestern side (Red River basin) basically is greater than that in the northeastern side (Chay River basin). In

which, some of the northern segment uplifted is greater than southern segment (Figure 5). This result is consistent with Le et al. (2001, 2004). In the northeastern side, where the Chay River fault cuts across at the foot of the slope, almost of basins with hypsometric curve are concave curve, except the basin 32 and 33. This is consistent with previous studies that Chay River fault is right-lateral slip (Nguyen, 2002; Phan et al., 2004, 2012; Ngo et al., 2006, 2011).

Table 1. The statistical moments of the hypsometric curve in the DNCV area (HI - Hypsometric integral, SK - Skew; KUR - Kurtosis, DSK - Density skew and DKUR - Density kurtosis)

No	HI	SK	KUR	DSK	DKUR	No	HI	SK	KUR	DSK	DKUR
1	0.335	0.526	2.141	0.666	2.002	23	0.188	1.174	3.626	1.402	4.109
2	0.294	0.451	2.016	0.614	1.758	24	0.255	0.816	2.726	0.839	2.485
3	0.294	0.452	2.003	0.736	1.956	25	0.214	0.953	2.900	1.189	3.158
4	0.269	0.487	2.055	0.658	1.829	26	0.190	1.169	3.341	1.550	4.203
5	0.272	0.609	2.164	0.964	2.370	27	0.169	0.848	2.560	1.070	2.653
6	0.309	0.595	2.221	0.612	1.804	28	0.128	1.183	3.328	1.499	3.987
7	0.250	0.579	2.200	0.555	1.724	29	0.156	1.019	2.766	1.346	3.282
8	0.284	0.788	2.607	0.867	2.410	30	0.137	0.626	2.138	0.904	2.147
9	0.305	0.598	2.236	0.662	1.916	31	0.205	0.983	2.974	1.218	3.249
10	0.311	0.643	2.285	0.752	2.039	32	0.344	0.550	2.232	0.327	1.591
11	0.329	0.667	2.482	0.444	1.864	33	0.290	0.463	1.999	0.495	1.525
12	0.320	0.642	2.386	0.517	1.848	34	0.254	0.752	2.498	0.841	2.277
13	0.370	0.605	2.375	0.339	1.759	35	0.220	0.860	2.570	1.063	2.644
14	0.333	0.706	2.562	0.525	2.010	36	0.227	1.033	3.112	1.324	3.638
15	0.270	0.848	2.792	0.873	2.539	37	0.261	0.825	2.673	0.970	2.661
16	0.329	0.717	2.523	0.649	2.070	38	0.252	0.727	2.456	0.780	2.155
17	0.347	0.773	2.667	0.714	2.292	39	0.214	0.885	2.812	0.976	2.689
18	0.304	0.833	2.877	0.688	2.366	40	0.185	1.137	3.381	1.415	3.916
19	0.259	0.994	3.176	1.082	3.150	41	0.240	0.933	3.032	0.970	2.916
20	0.282	1.126	3.756	1.003	3.402	42	0.257	0.889	3.000	0.820	2.665
21	0.213	1.302	4.040	1.575	4.805	43	0.199	1.306	4.100	1.563	4.875
22	0.239	1.106	3.679	1.106	3.567	44	0.191	1.064	3.101	1.453	3.925

According to Al Hamdouni et al. (2008), the hypsometric curve often has convex curve when HI index greater than 0.5; intermediate form between the concave and convex shape (S-shape) or "straight-shape" when the HI value in the range of 0.4 to 0.5 and the HI-value less than 0.4, the hypsometric curve has a concave shape. In the study area, as the Table 1, Figure 5 and Figures 6a, b, the HI

values of the hypsometric curve with straight-shape and S-shape are less than 0.4 and smallest is 0.25. Thus, when using and analyzing the HI index in different areas, need to combine with its hypsometric curve. Because in many cases, the basins with similar hypsometric integrals but different shapes (Pérez-Peña et al., 2009). In that cases, these other statistical moments are necessary

to consider for the hypsometric analysis (Figure 7a, b).

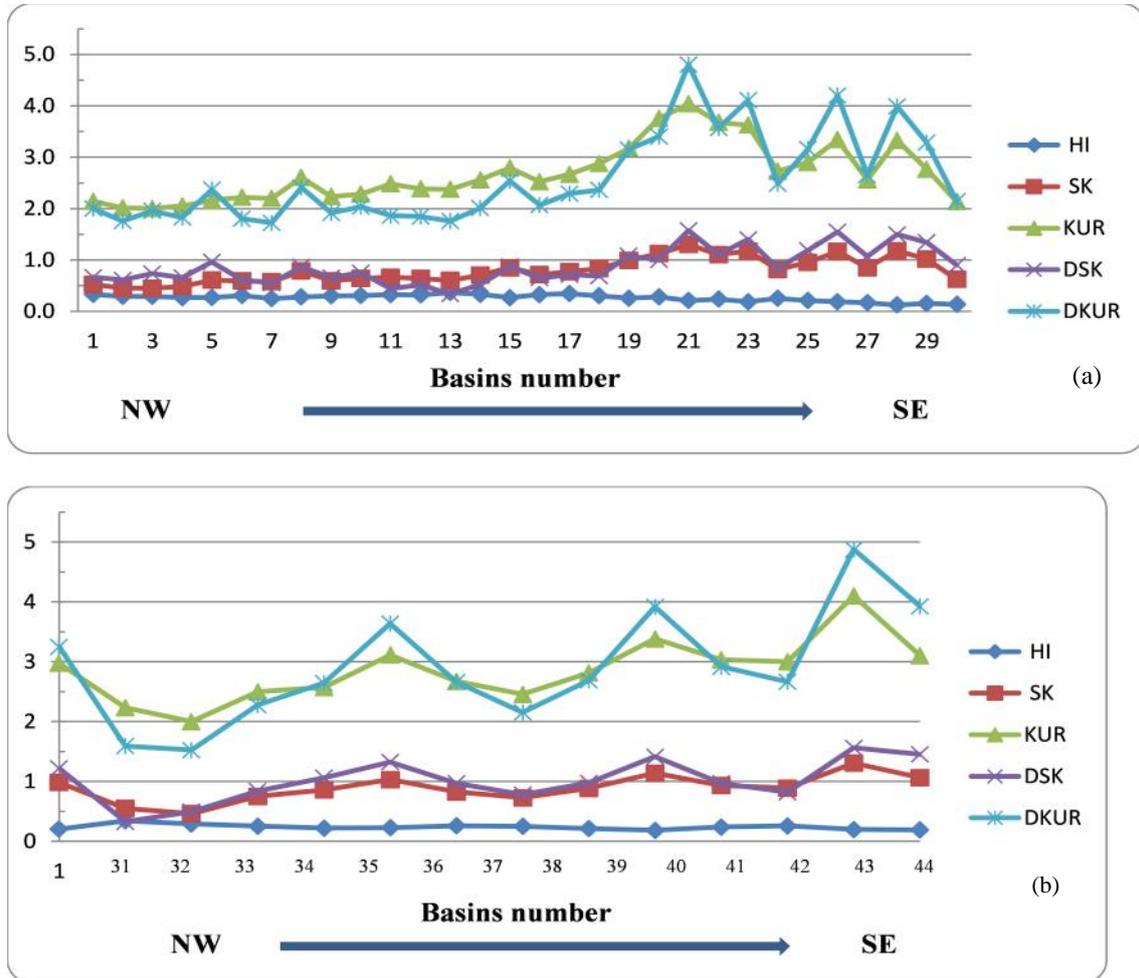


Figure 7. a) The variation of the statistical moments of the hypsometric curve (the basins in the southwestern side of the DNCV); b) The variation of the statistical moments of the hypsometric curve (the basins in the northeastern side of the DNCV)

In the study area, according to the results in Table 1, Figure 7 (a, b) and in the direction from northwest to southeast, unless the HI index (downward trend), basically, other statistical moments of the hypsometric curve are likely to increase in both sub-basin systems of the Red River and Chay River. As the results of skew, the values range from 0.45 to 1.3 ($SK > 0$), this mean the basins in the study with the geomorphological processes almost are represent the amount of headward erosion in

the upper reach of the basin (Harlin, 1978; Luo, 2000; Pérez-Peña et al., 2009). This trend is basically increased in the direction from northwest to southeast to the basin of the study area. Consistent with SK index, the DSK index also reflects a larger slope in the upper reach of the basin and also showed upward trend from northwest basins to the southeast basins. In addition, a larger value of kurtosis (almost is the concave shape in the southern of the DNCV in both the northeast

and southwest sides; Table 1 and Figure 7) signifies erosion on both upper and lower reaches of a basin (Harlin, 1978; Luo, 2000). These results also showed, in the basins which have large KUR values then so are DKUR values (Table 1 and Figure 7). This means that, this basin also has large slope in the middle part of the basin (Harlin, 1978; Luo, 2000; Pérez-Peña et al., 2009). What makes this area contain these features (erosion process in the both of the upstream and downstream area, and addition large slope in the middle part)? According to Le (2001, 2004) and Ngo (2011), the regional topography has stepped clearly. This step by the heterogeneously raising activities and the active fault branches (of the Red River and the Song Chay Faults) on both the northeast and southwest sides of the DNCV area (Le, 2001, 2004; Ngo, 2011).

According to Pérez-Peña et al. (2009), the value increases of the KUR and DKUR indexes (when the same hypsometric curve and hypsometric integral index) often show upward trend of recent tectonic activity. However, in the study area, the higher anomalies of the KUR and DKUR indexes still lie in the basins with the hypsometric curve showing concave shape (Table 1, Figure 5 - Figure 7). Thus, if only individual basins with curve concave shape are considered, basins with larger KUR and DKUR values will show stronger tectonic activity. If all basins of the study area are considered, the higher anomalies KUR and DKUR indexes at the basins with concave curve shape possibly suggest the following remarks: According to basic hypsometric curve model and geomorphological development cycles shown by the changes of the hypsometric curves (Figure 3), the hypsometric curves with concave curve shape are the oldest stage of geomorphological cycles, and it is the final stage of the cycle to stabilize the tectonic cycle to pass on to a new tectonic activity cycle. This means, the tectonic active in the study area possibly is the last period of

stabilization tectonic cycle and the beginning of a new tectonic activity cycle. If so, the above assumption is appropriated in anticipation of Allen (1984) to repeat the cycle of large earthquakes along the Red River Fault Zone is about 1800 years, while in the region of Yunnan, China had strong earthquake occurred 8.1 to 8.3 on the Richter scale and occurred approximately 1000 to 2000 years ago.

In summary, the hypsometric curves in the study area are mainly concave shapes; some curves are intermediate form between the concave and convex shape (in "straight" and "S" shape). HI index is basically small and tends to decrease to the southeast. The skew and kurtosis and their density function are basically large and increasing trend to the southeast. An overview, the recent tectonic activity (uplift - lower) in the study area is generally weak. In which, the southwestern side is being lifted higher than the north-eastern side. The northern part is being lifted larger than the southern part. In the region and surrounding area, the strong uplift activities and increased gradually in the Pliocene-Quaternary (modeled after Le et al., 2004) could have stopped at certain time in the past. The current geomorphic processes are mainly headward erosion in the upstream. These results will be clarified in the next study when there is a combination of many different geomorphic indices.

6. Conclusions

The hypsometric curves and its statistical moments are useful tools to assess the geomorphological processes and recent tectonic activity of the region as well as the comparison between different zones.

The Day Nui Con Voi area has revealed 3 curves such as "straight- shape", "S- shape", and concave curves. The concave curve is the most common widely distributed in the northeast side and the southern part of the southwestern side of the DNCV area. The hypsometric integral (HI) values are rather

small, the largest value is 0.37 whereas the smallest one is 0.128. Other statistical moments of the hypsometric curve i.e. skew (SK), kurtosis (KUR), and the density function (density skew - DSK and density kurtosis-DKUR) have great values and increase in the south direction of the area study.

The recent active tectonic activities (uplift-lower) of the study area are generally weak. However, they are also not completely homogeneous and can be distinguished by different levels. The southwestern side is being lifted higher than the north-eastern side. The northern part is being lifted larger than the southern part. In the region, the uplift activities were increased gradually in the Pliocene-Quaternary and could have stopped at certain time in the past. The current geomorphic processes are mainly headward erosion in the upstream.

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